

1 **A global-scale intercomparison of Triple Collocation Analysis- and**
2 **ground-based soil moisture time-variant errors derived from**
3 **different rescaling techniques**

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10 **Abstract:** Accurate specification of spatiotemporal errors of remotely sensed soil
11 moisture (SM) data is essential for a correct assessment of their utility and optimally
12 integrating multiple SM products or assimilating them into hydrological models.
13 Although Triple Collocation Analysis (TCA) has been widely used to provide SM
14 errors, the impact of rescaling technique on the TCA error estimates has not received
15 major attention, which can lead to biased and inaccurate error estimates. Moreover,
16 current knowledge about time-variant SM errors derived from TCA is still very
17 limited, which hampers the advance of applying time-variant errors in data merging
18 and data assimilation studies efficiently. Based on these considerations, this work
19 aims to advance the use of the TCA for characterizing errors with a focus on the
20 rescaling techniques, and validating TCA-based time-variant errors using global
21 ground measurements in 759 grid cells. To this end, the Advanced Scatterometer
22 (ASCAT) and four passive-based SM products, including Soil Moisture and Ocean
23 Salinity Level-3 (SMOSL3), SMOS INRA-CESBIO (SMOSIC), Soil Moisture Active
24 Passive Level-3 (SMAPL3), and SMAP INRAE BORDEAUX (SMAPIB) SM
25 products were considered. The time-variant error term here denotes an aggregate error
26 magnitude over a 101-day moving-time-window. It is found that different selection of
27 the rescaling technique considered in TCA led to TCA error estimates with
28 significantly different accuracy when ground-based errors are regarded as the
29 benchmark. The optimal combination strategy to implement TCA is applying TCA to
30 SM anomalies and rescaling the errors by coefficients derived from the TCA model.
31 Pearson's correlation with ground-based time-variant errors is 0.62, 0.72, 0.83, 0.89,

32 and 0.93 for SMAPIB, SMAPL3, SMOSIC, SMOSL3, and ASCAT SM, respectively.

33 Considering time-variant errors in applications is necessary since time-variant errors

34 deviate from time-invariant errors by 50% when errors are rescaled by TCA model

35 parameters. Time-invariant errors are greater than time-variant errors when SM

36 products are rescaled against a reference dataset while the opposite conclusion can be

37 drawn when errors are rescaled by the TCA coefficients. TCA- and ground-based

38 methods provide consistent evaluations in 74.7% (77.3%), 75.8% (79.8%), 79.6%

39 (81.1%), and 78.6% (79.7%) of the analysis period on average (median) for the TCA

40 implementations with SMAPL3, SMAPIB, SMOSL3, and SMOSIC SM, respectively.

41 The error analysis reveals that TCA typically underestimated ASCAT errors while

42 overestimated passive SM errors when considering ground-based evaluation as the

43 benchmark. Moreover, TCA was found to have relatively less power to efficiently

44 characterize SM errors in croplands when compared with other land cover types. This

45 study validated TCA time-variant errors using ground measurements and compared

46 TCA- and ground-based evaluation performances on a global scale. Our work arouses

47 particular attention to the rescaling technique selection considered in TCA, which is

48 crucial for accurately characterizing SM errors and efficiently using them in various

49 hydrometeorological applications.

50 **Keywords:** soil moisture; rescaling technique; Triple Collocation Analysis; time-

51 variant error

52

53 **1 Introduction**

54 Soil moisture (SM) plays an important role in modeling hydrological processes
55 such as runoff and evapotranspiration, and links water, energy, and carbon cycles
56 (Jackson, 1993; Houser et al., 1998; Western et al., 2002; Daly and Porporato, 2005).
57 SM data can be applied in many disciplines such as drought monitoring, flood
58 prediction, crop productivity forecasts, irrigation planning, and weather forecasting
59 (Sims et al., 2002; Narasimhan and Srinivasan, 2005; Bolten et al., 2010; Wanders et
60 al., 2014). SM observations can be obtained from ground measurements, hydrological
61 modeling, and satellite observations, and each of them has its distinctive error
62 characteristic. Learning about the error characteristics of various SM products is
63 important as it significantly influences the uncertainties of the hydrological models
64 driven by the SM content and therefore has a great impact on the interpretation of the
65 model simulation. Furthermore, stochastic data assimilation relies on accurate error
66 specifications for the observations and model predictions (Reichle, 2008). Data fusion
67 studies also require accurate error specifications (Crow et al., 2015; Gruber et al.,
68 2017) or signal-to-noise ratio information (Kim et al., 2022) to optimally integrate
69 multiple SM products. An accurate specification of SM error variability in space
70 (such as Gruber et al., 2015 and Wu et al., 2018) and in time (such as Loew and
71 Schlenz, 2011; Zwieback et al., 2012; Su et al., 2014a; Wu et al., 2021) may lead to
72 further improvements for data assimilation and data fusion studies.

73 Triple Collocation Analysis (TCA) is a popular evaluation method to provide
74 relative errors for SM products derived from different platforms without requiring an

75 absolute ‘truth’ (Dorigo et al., 2010). It was initially proposed by Stoffelen (1998) to
76 solve the issue of error estimation for sea wind and wave height and was introduced to
77 provide SM error estimates by Scipal et al. in 2008a. Besides, SM uncertainty can be
78 obtained by a three-corned hat method, which was recently applied by Liu et al. (2021)
79 to evaluate eleven SM products in the Qinghai-Tibet Plateau. Many studies have
80 applied TCA to evaluate remotely sensed SM retrievals (such as Leroux et al., 2013;
81 Su et al., 2014a; Su et al., 2014b; Chakravorty et al., 2016; Kim et al., 2018), model
82 simulations (such as Dorigo et al., 2010; Al-Yaari et al., 2014), reanalysis SM (such
83 as Scipal et al., 2008a; Scipal et al., 2010; Yilmaz and Crow 2014; Miyaoka et al.,
84 2017), or ground-based SM measurements (Miralles et al., 2010; Chen et al., 2018;
85 Wu et al., 2021). Currently, TCA-based SM error estimates have already been widely
86 used in SM validation (Dorigo et al., 2010; Miyaoka et al., 2017; Kim et al., 2021),
87 data assimilation (Gruber et al., 2019), and data fusion (Gruber et al., 2017; Peng et
88 al., 2021) studies.

89 TCA has been typically applied to provide time-invariant errors in the whole
90 investigation period. However, SM retrieval errors are known to vary with time due to
91 vegetation phenology, changes in surface roughness, and variable environmental
92 conditions (Ulaby et al., 1983; Famiglietti et al., 2008; Zwieback et al., 2018). To
93 account for the time-variant feature of SM products, Loew and Schlenz (2011),
94 Zwieback et al. (2012), Su et al. (2014a), and Wu et al. (2021) proposed various time-
95 window-based TCA schemes to estimate time-variant errors at different timescales by
96 relaxing the stationary assumption underlying the TCA. Several studies tried to

97 consider and include time-variant error characteristics in applications. For example,
98 Khan et al. (2014) proposed a multi-model data merging approach that considers
99 monthly covariance matrix of the forecast errors, which was successfully applied to
100 the sea surface temperature forecasts and achieved better performance than time-
101 invariant forecasts. In addition, Kim et al. (2016) combined multiple satellite-based
102 SM datasets by taking time-variant weights into account and reported this method
103 outperforms the time-invariant approach. Therefore, an extension of TCA to temporal
104 domain may greatly benefit many applications. Nevertheless, accuracy of the time-
105 variant errors derived from window-based TCA scheme is not well known by the SM
106 community, which should be investigated before applying time-variant SM errors in
107 actual applications.

108 As increasing number of applications use TCA to characterize errors for large
109 spatial and temporal datasets, the reliability of TCA has drawn attention from several
110 fields, such as soil freeze/thaw (Li et al., 2022), root zone SM (Xu et al., 2021), and
111 satellite-based surface albedo (Wu et al., 2019). However, more work is still needed
112 to verify TCA-based SM spatiotemporal errors using ground measurements as such
113 validation on a global scale has not been fully investigated in previous studies. We
114 argue that this validation is important as it helps researchers learn about the accuracy
115 and reliability of the TCA, and better interprets the TCA error estimates. Previous
116 work on this topic had mostly a regional focus or did not consider time-variant errors.
117 For example, Brocca et al. (2011) and Chen et al. (2016) reported that TCA- and
118 ground-based estimates, i.e., error and satellite-versus-truth correlation coefficient, are

119 strongly consistent in Europe and USA, respectively. By comparing the TCA- and
120 ground-based errors, Loew and Schlenz (2011) investigated the point-to-area
121 sampling error in Southern Germany. Notably, Kim et al. (2020) compared TCA-
122 based fMSE (fractional Mean-Square-Error) with the conventional fMSE derived
123 from ground measurements and showed these two kinds of methods yielded
124 consistent evaluation results. Similarly, Zhang et al. (2021) investigated the TCA- and
125 ground-based satellite-versus-truth correlation coefficients and reported that they
126 yielded consistent spatial distribution. However, studies by Kim et al. (2020) and
127 Zhang et al. (2021) only focus on the spatial domain, i.e., error metrics are assumed to
128 be constant over time. The performance of TCA in the temporal domain remains
129 unclear currently. Moreover, Yilmaz and Crow (2014) presented numerical and
130 analytical comparisons for ground- and TCA-based SM errors and found that ground-
131 based errors are often higher than TCA-based errors. A similar result was also
132 confirmed by Dorigo et al. (2015) who pointed out that TCA errors were consistently
133 lower than ground-based errors. Based on current studies, a global-scale
134 intercomparison of TCA- and ground-based SM errors is highly needed to advance
135 the TCA applications, especially for time-variant errors as they have great potential to
136 improve the performance of data merging and data assimilation systems.

137 To accurately estimate time-variant SM errors using the TCA method, two key
138 issues are typically ignored in current TCA studies, which may lead to inaccurate and
139 biased SM error estimates. The first issue is the rescaling technique selection
140 considered in TCA. The rescaling process is an essential step in TCA as it removes

141 relative differences between the considered datasets and makes the resulting error
142 estimates comparable for the statistical comparison purpose. The rescaling technique
143 applied in TCA not only can minimize the impact of representativeness errors derived
144 from different spatial resolution, spatiotemporal mis-alignment, and different vertical
145 measurement support (Gruber et al., 2013; Chen et al., 2017; Molero et al., 2018) but
146 also can implicitly compensate for different units used in the considered datasets
147 (Gruber et al., 2020). The rescaling techniques applied in TCA can be divided into
148 two categories: methods that rescale SM inputs against a selected reference dataset
149 prior to TCA implementation (Scipal et al., 2008a), and methods that rescale errors by
150 parameters derived from the TCA model after the TCA implementation (Gruber et al.,
151 2016a). There are several statistical techniques to rescale SM observations prior to
152 TCA, such as variance and mean matching (Dorigo et al., 2010; Miralles et al., 2010;
153 Kim et al., 2020), normalization (Fascetti et al., 2016; Pierdicca et al., 2015; Wu et al.,
154 2021), CDF (Cumulative Distribution Function) matching (Doubkova et al., 2012;
155 Gruber et al., 2014; An et al., 2016; Zhuang et al., 2020), and linear regression
156 method (Scipal et al., 2008a). Or normalizing all datasets into a common observation
157 space prior to TCA by calculating their z-scores. Generally, both these two kinds of
158 rescaling techniques can address the first-order (additive) biases for the considered
159 datasets. However, Gruber et al. (2020) reported that the statistical matching cannot
160 address the second-order (multiplicative) biases while the method that rescales errors
161 using the TCA model parameters can. Most current studies ignored the impact of
162 second-order (multiplicative) biases on validation and error characterization. Such

163 impact of second-order (multiplicative) biases resulted from the rescaling technique
164 selection on the TCA error characterization remain unclear and needs further
165 investigation (Gruber et al., 2020). Nevertheless, selection of the rescaling technique
166 has not received major attention in current TCA studies. Recently, both Kim et al.
167 (2020) and Gruber et al. (2020) pointed out the necessity to investigate the impact of
168 rescaling technique selection on the final TCA error estimates.

169 The other important issue that influences TCA is the selection of SM inputs.
170 There are two options for the SM inputs. One uses the original SM measurements
171 directly (such as Dorigo et al., 2015; Gruber et al., 2017; Wu et al., 2018; Wu et al.,
172 2021) while the other one uses SM anomalies (such as Dorigo et al., 2010; Draper et
173 al., 2013; Su et al., 2014a; Miralles et al., 2010; Chakravorty et al., 2016; Kim et al.,
174 2020). In practice, a popular technique to obtain short-term anomalies relies on
175 moving-window-based averages over the investigation period, carried out by Miralles
176 et al. (2010), Crow et al. (2012), and Draper et al. (2013). The anomaly-based
177 approach removes the seasonal effects underlying SM time-series that can artificially
178 enhance the correlations between two SM time-series (Scipal et al., 2005; Scipal et al.,
179 2008b) and therefore, reveals the ability of the SM products to capture the short-term
180 events of drying and wetting (Dorigo et al., 2010; Al-Yaari et al., 2014). Working on
181 the original time series also makes sense as it captures other properties of the data sets,
182 even though the errors cross-correlations may be higher (Miralles et al., 2010; Draper
183 et al., 2013). These two methods are complementary as both are needed to describe
184 the quality of the SM data. However, no study to date makes a quantitative

185 comparison for the final TCA error estimates derived from these two input selections.
186 If TCA errors have small differences for the two methods, absolute-based approach is
187 more favorable given its simplicity in practice.

188 Based on the considerations above, we first explored the impact of temporal
189 interpolation on the accuracy of TCA error estimates in Section 3.1. Second, the
190 impacts of rescaling technique and SM inputs selection on the accuracy of TCA time-
191 invariant and time-variant errors were investigated in Section 3.2. Correlations and
192 RMSE (Root Mean Squared Error) between TCA- and ground-based time-variant
193 errors were also compared for the ASCAT and four passive-based SM data, i.e.,
194 SMOSL3, SMOSIC, SMAPL3, and SMAPIB SM products, in six land cover types.
195 Third, the relative difference between time-invariant and time-variant errors was
196 quantified in Section 3.3. Then, the evaluation consistency between TCA- and
197 ground-based methods is explored in Section 3.4. Finally, time-variant errors were
198 compared for ASCAT and the four passive SM products based on the TCA- and
199 ground-based methods in Section 3.5, along with TCA- vs. ground-based evaluations
200 over six land cover types.

201 **2 Data and Methodology**

202 This section introduces SM products and methods used. Sections 2.1 and 2.2
203 provide information about Soil Moisture and Ocean Salinity (SMOS), Soil Moisture
204 Active Passive (SMAP), ASCAT SWI (Soil Water Index), Global Land Data
205 Assimilation System Version 2.1 (GLDAS2), and ERA-Interim SM products, along
206 with International Soil Moisture Network (ISMN) ground measurements. After briefly

207 describing TCA and Quadruple Collocation Analysis (QCA) methods in Section 2.3,
208 Section 2.4 details four rescaling techniques used in TCA implementation.
209 Conventional ground-based method to estimate SM errors is presented in Section 2.5
210 and the approach to test the robustness of temporal interpolation applied in the TCA
211 time-variant scheme is described in Section 2.6. Finally, Section 2.7 defines the
212 evaluation metrics used in our work: Overall Relative Difference (ORD), Relative
213 Difference (RD), RMSE, and Root Mean Squared Difference (RMSD).

214 **2.1 SMOS, SMAP, ASCAT, GLDAS2, and ERA-Interim SM products**

215 The Soil Moisture and Ocean Salinity (SMOS) mission is designed to observe
216 SM content in the global landmass surface (0-5 cm) and salinity over the oceans based
217 on multi-angular brightness temperature data observed by a radiometer operating at L-
218 band (Kerr et al., 2012). It was launched by European Space Agency (ESA) in
219 November 2009 and contributes to improving our understanding of the global water
220 cycle and weather/seasonal climate forecasting. SMOS collects brightness
221 temperature data at a local overpass time of 06:00 PM for descending pass and 06:00
222 AM for ascending pass, respectively. Here, the SMOS Level-3 (SMOSL3) SM
223 product with version 3.3 was used (Al Bitar et al., 2017). This product was developed
224 by CATDS-PDC (Centre Aval de Traitement des Données SMOS - Production &
225 Dissemination Center) and provides SM observations projected on Equal-Area
226 Scalable Earth Grid (Version 2) with a spatial resolution of 40 km.

227 Data screening for SMOSL3 SM was based on the DQX (data quality index)
228 affiliated with the CATDS-PDC SMOSL3 product and SM observations were filtered

229 out in our work when its associated DQX value is greater than 0.007 (Chen et al.,
230 2018; Wang et al., 2021). The DQX value quantifies the error in SM observations in
231 volumetric SM units (Al-Yaari et al., 2014). The RFI (Radio Frequency Interference)
232 issue was not taken into consideration for SMOSL3 data screening as we expect more
233 SM observations participated in the time-variant error estimation, especially in
234 regions such as Eurasian.

235 Besides SMOSL3, SMOS INRA-CESBIO (SMOSIC) SM product of version 2
236 was also considered in our work. The SMOSIC SM product was constructed by INRA
237 (Institute National de la Recherche Agronomique) and provides SM values derived
238 from the SMOS L3 brightness temperature data with a spatial resolution of 25 km.
239 One of the main characteristics of the SMOSIC SM is the maximal independence of
240 auxiliary data during its retrieval process, which is different from the SMOSL3 SM
241 that strongly relies on auxiliary data (Wigneron et al., 2021). The SMOSIC SM
242 product was used here to increase the robustness of our TCA conclusions since the
243 SMOSL3 SM product is not independent with model-based or reanalysis SM data and
244 therefore may hinder the independence assumption required by the TCA and may
245 have potential influence on the final intercomparison and validation. To guarantee the
246 data quality of SMOSIC SM data, only SM values associated with scene flags ≤ 1
247 were retained in the following analyses. The scene flags were affiliated with the
248 SMOSIC product and can be used as an indicator to detect events that influence the
249 SMOS L3 brightness temperature observations used in the retrieval process, such as
250 strong topography, frozen condition, and water body contamination. Similar to the

251 data filtering of the SMOSL3, the RFI problem was not considered for the SMOSIC
252 SM in our analysis. Readers are recommended to refer to Wigneron et al. (2021) to
253 find out more information about the SMOSIC SM product.

254 The SMOSL3 and SMOSIC SM datasets were reconstructed from local time-
255 based into UTC time-based to match the ASCAT, GLDAS2, and ISMN ground
256 measurements. This reconstruction was completed by considering the navigational
257 time zones based on longitude values and local statutory deviation was not considered
258 in such transformation (Kim et al., 2018). After the transformation, SM values at
259 ascending and descending orbits were averaged when they are found on the same day.
260 Consequently, SMOSL3 and SMOSIC SM daily datasets with UTC stamps were
261 constructed and used here. Readers can find the conceptual map (Fig. S1) of UTC
262 zones for converting the local time, i.e., ascending for 06:00 AM and descending for
263 06:00 PM, of SMOS dataset to match other SM data.

264 The Soil Moisture Active Passive (SMAP) satellite is designed by NASA
265 (National Aeronautics and Space Administration) to provide global surface (0-5 cm)
266 SM observations with a spatial resolution of 40 km (Entekhabi et al., 2010). Since a
267 malfunction was found in the SMAP radar system, only the L-band radiometer
268 onboard SMAP satellite collects observations at a local overpass time of 06:00 AM
269 for descending pass and 06:00 PM for ascending pass, respectively (Wu et al., 2020).
270 The SMAP Level-3 SM product in version 8 (SMAPL3) was used here (O'Neill et al.,
271 2021). Begin with this version, the Dual Channel Algorithm (DCA) is applied as a
272 new baseline algorithm in the retrieval process, which departs from prior versions that

273 used the Single Channel Algorithm-Vertical Polarization (SCA-V) as the baseline
274 algorithm. Data screening for the SMAPL3 SM product was implemented using the
275 retrieval quality flag affiliated with the SMAPL3 product. Only SM observations with
276 recommended quality were considered in our analysis. The quality flag indicates
277 whether unfavorable environmental conditions, such as frozen soil, snow cover, flood,
278 steeply sloped topography, urban area, and dense vegetation, occurred during the SM
279 retrievals.

280 Besides SMAPL3, a recently developed SM product by INRAE Bordeaux (Li et
281 al., 2022), i.e., SMAP INRAE BORDEAUX (SMAPIB) SM data, was also
282 considered in our analysis. The SMAPIB SM product applied a new 2-Parameter
283 retrieval algorithm, i.e., SM and vegetation optical depth, to the dual-polarized
284 brightness temperature observations collected from SMAP satellite during its SM
285 retrieval process. Currently, only SM observations acquired from the descending
286 (06:00, local overpass time) orbits are available and can be freely accessed from
287 <https://ib.remote-sensing.inrae.fr/>. Similar to the SMOSIC data, this SM product was
288 used here as it does not use any modeled or reanalysis SM data as input in its retrieval
289 algorithm, which makes it a favorable SM dataset to be applied to the TCA method
290 that requires a strong independence assumption. As suggested by Li et al. (2022), only
291 SM observations with scene flag value ≤ 1 were kept in the following analyses as
292 corresponding SM values are less affected by the frozen soil, strong topography, and
293 water body contamination. To match with other SM data, the SMAPL3 and SMAPIB
294 SM datasets were reconstructed from local time-based to UTC time-based using the

295 same method applied in the SMOSL3 and SMOSIC data preprocessing.

296 The Advanced Scatterometer (ASCAT) is a C-band (5.2 GHz) active sensor
297 onboard the MetOp satellite (Bartalis et al., 2007). The ASCAT SWI (Soil Water
298 Index) product (Wagner et al., 1999) with T=1 was used here, which describes SM
299 content in the top soil layer (Albergel et al., 2008; Paulik et al., 2014). This product
300 was obtained from CGLS (Copernicus Global Land Service,
301 <https://land.copernicus.eu/global/products/swi>), which provides SM observations
302 projected on a regular latitude/longitude grid with a spatial resolution of $0.1^\circ \times 0.1^\circ$
303 in units of percentage. The ASCAT SWI is a daily product with a time stamp at 12:00
304 UTC regardless of the actual observation time (Paulik et al., 2014). Previous studies
305 have validated this SM product using remotely sensed SM or ground measurements,
306 such as Albergel et al. (2009), Brocca et al. (2010), Brocca et al. (2011), and Paulik et
307 al. (2014).

308 The ASCAT SWI values were masked out in our work when its surface state flag
309 indicates frozen soil conditions. The surface state flag contains three different soil
310 conditions, namely frozen, unfrozen, and melting states of the soil surface based on
311 the temperature observations. Moreover, the SWI product provides a quality flag
312 (QFLAG) that describes the number of available ASCAT SM observations considered
313 during the SWI calculation. Here, SWI values were filtered out when the
314 corresponding QFLAG value is smaller than 10% as these SWI values become
315 unreliable due to limited ASCAT SM inputs (Wu et al., 2021).

316 SM product simulated from the GLDAS2 was used as a third independent

317 estimate of SM applied in our TCA. GLDAS2 SM product was created by the Noah
318 land surface model that uses a combination of model and observation data (excluding
319 SM products) as its forcing data (Rodell et al., 2004). Only SM simulations in the top
320 soil layer (0-10 cm) were used here since ASCAT and SMOS/SMAP SM
321 observations only represent SM content of the topsoil layer in several centimeters.
322 The GLDAS2 provides 3-hourly SM simulations projected on a regular
323 latitude/longitude grid of $0.25^{\circ} \times 0.25^{\circ}$. Here, a daily GLDAS2 SM product was
324 constructed by averaging all available SM simulations within the same day. To
325 exclude SM values simulated under frozen soil conditions, GLDAS2 SM values were
326 screened if the corresponding GLDAS2 soil temperature value is lower than 0°C .

327 The ERA-Interim SM product is a global reanalysis dataset derived from an
328 advanced data assimilation system (Balsamo et al., 2015). It was used as an
329 independent SM here to test the spatial representativeness of the ground
330 measurements following Dorigo et al. (2015). This dataset is projected on a regular
331 latitude/longitude grid with a spatial resolution of $0.25^{\circ} \times 0.25^{\circ}$ and provides SM
332 analysis values with UTC stamps in the four soil layers, i.e., 0-7 cm, 7-28 cm, 28-100
333 cm, and 100-289 cm. Here, only SM data in the soil layer of 0-7 cm were considered.
334 A daily ERA-Interim SM dataset was reconstructed by averaging all available SM
335 values within the same day. Given that the ERA-Interim product is only available
336 until 31 August, 2019, ERA-Interim SM data from July 3, 2012 to 31 August, 2019
337 were used in our analysis. The ERA-Interim SM values were screened when the
338 associated GLDAS2 soil temperature value is lower than 0°C to exclude SM values

339 simulated under frozen soil conditions.

340 In our work, the projection grid of the GLDAS2 SM product was selected as the
 341 reference projection. ASCAT, SMOS, SMAP, and ERA-Interim SM observations
 342 were resampled on the GLDAS2 grid using the nearest neighbor interpolation. Since
 343 multiple SM products were considered here, their different overpass time stamps were
 344 accounted for by binning all SM products to a regularized time step (Gruber et al.,
 345 2020), i.e., daily UTC time here. A summary of the SM datasets used in our work is
 346 provided in Table 1. The timeframe of our analysis starts from July 3, 2012 to August
 347 31, 2020 with a total of 2982 days. Although the discrepancy in measurement depth is
 348 considerable for ASCAT, SMOS, SMAP, GLDAS2, and ERA-Interim, Albergel et al.
 349 (2008) and Brocca et al. (2011) pointed out that SM in the upper 10 cm is strongly
 350 correlated with surface SM (e.g., 0-5 cm).

351 Table 1 A summary of the SM products used in this work

SM products	Spatial resolution	Temporal resolution	Nominal observation depth	Temporal coverage	Version	Projection grid
GLDAS2	0.25°×0.25°	3-hourly, UTC	0-10 cm	January 2000 to present	2.1	regular lat/lon grid
ASCAT SWI	0.1°×0.1°	daily at 12:00, UTC	0-5 cm	January 2007 to present	3.0	regular lat/lon grid
SMOSL3	40 km	6 AM and 6 PM local time	0-5 cm	January 2010 to present	3.3	EASE-Grid 2.0
SMOSIC	25 km	6 AM and 6 PM local time	0-5 cm	January 2010 to present	2.0	EASE-Grid 2.0
ERA-Interim	0.25°×0.25°	3-hourly, UTC	0-7 cm	January 1979 to August 2019	2.0	regular lat/lon grid
SMAPL3	36 km	6 AM and 6 PM local time	0-5 cm	March 2015 to present	8	EASE-Grid 2.0
SMAPIB	36 km	6 AM local time	0-5 cm	March 2015 to present	1.0	EASE-Grid 2.0

352 Note: Equal-Area Scalable Earth Grid, Version 2.0 (EASE-Grid 2.0)

353 In addition, SMOS and SMAP SM data were temporally interpolated here by

354 filling the gaps existed in their time series with the average of neighboring SM values
355 included in a 3-day time window. This will significantly increase the statistical power
356 of the TCA method since more SM triplet samples can participate in the TCA
357 calculation (Leroux et al., 2013; Wu et al., 2021). Limited samples, e.g., 30 samples,
358 can lead to many error estimates that are not convergent, which reduces the robustness
359 of the moving-window-based TCA method. However, the temporal interpolation
360 influences the final TCA errors inevitably due to the additional interpolation errors
361 and such impact should be investigated before applying the TCA time-variant scheme
362 to the temporally interpolated SM datasets (detailed in Sec. 2.6 and Sec. 3.1).

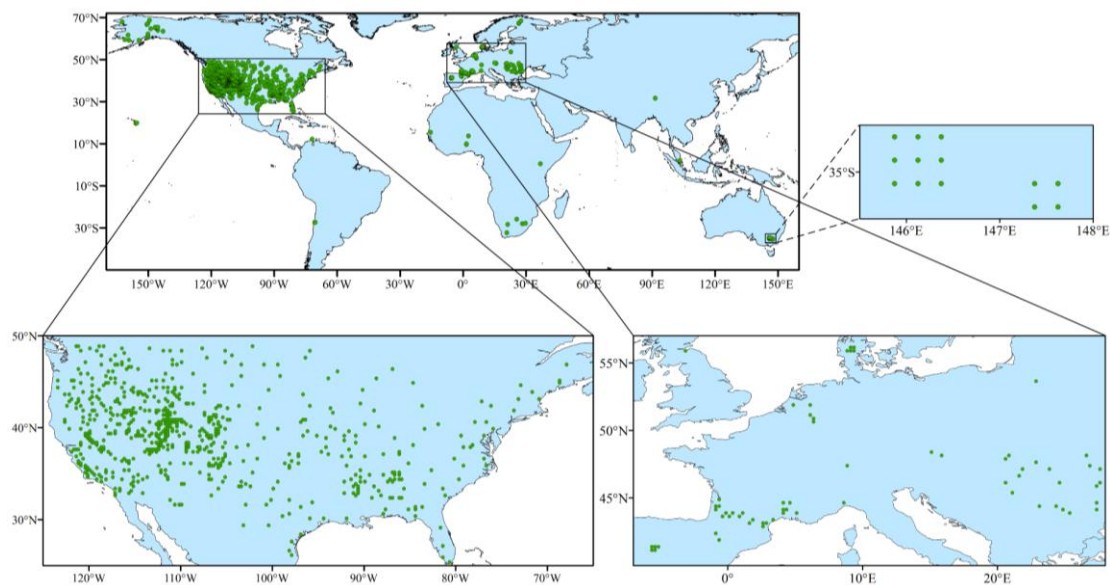
363 **2.2 ISMN ground measurements**

364 Ground-based SM measurements are often considered to be the most accurate
365 representation of the true SM content even though their spatial support is very small.
366 The ISMN ground measurements from 759 stations belonging to 25 sparse networks
367 (Dorigo et al., 2011) were used here to validate the TCA evaluation performance. To
368 be consistent with the UTC time of other SM datasets, we took an average of all
369 available SM values on the same day.

370 Stations were selected using several criteria. First, only stations of which the
371 measurement interval was restricted within the upper 10 cm of the top soil layer were
372 considered. Station with the shallowest depth was selected when stations with
373 multiple depths are available in the same location (e.g., one taken at 5 cm and one
374 taken at 10 cm). Second, only stations where the number of available SM values is
375 greater than 500 were selected to guarantee sufficient data samples applied in the

376 TCA. Third, an areal representative station (Dorigo et al., 2015) was selected using a
377 3rd independent dataset, i.e., the ERA-Interim SM product, if more than one station
378 was found in the experimental grid cell. Correlation coefficients between ground
379 measurements and ERA-Interim, GLDAS2, SMOS/SMAP, and ASCAT SM data
380 were calculated and the station with the highest correlation average was selected in
381 the following analysis. Finally, ISMN ground measurements were screened by the
382 ISMN quality flags (Dorigo et al., 2013) to guarantee their data quality. After filtering
383 ISMN stations in the top soil, 759 stations remained available in our work. Fig. 1
384 shows the 759 experimental pixels that include these selected ISMN stations and
385 Table 2 summarizes the ISMN sparse networks, number of stations used, and
386 associated references.

387 Referring to the classification scheme of the MODIS IGBP dataset (Friedl et al.,
388 2002), land cover types associated with the experimental grid cells mainly include
389 grasslands (47.4%), croplands (15.0%), woody savannas (12.1%), savannas (8.8%),
390 forests (7.5%), and open shrublands (4.9%). These six land cover types are used to
391 categorize our results in the following analyses.



392

393 Fig. 1 Spatial distribution of 759 grid cells that include selected ISMN stations.

394

395

Table 2 ISMN ground stations from sparse networks

Network	Number of stations	Reference
AMMA-CATCH	3	Pellarin et al. (2009), Mougin et al. (2009), Cappelaere et al. (2009), de Rosnay et al. (2009)
ARM	4	http://www.arm.gov/
BIEBRZA_S-1	1	http://www.igik.edu.pl/en
COSMOS	5	Zreda et al. (2012)
CTP_SMTMN	1	Yang et al. (2013)
DAHRA	1	Tagesson et al. (2015)
FMI	3	http://fmiarc.fmi.fi/
FR_Aqui	3	Institute of Agricultural Research
GROW	7	https://growobservatory.org/index.html
HOBE	5	Bircher et al. (2012)
IMA_CAN1	1	Biddoccu et al. (2016)
IPE	2	Instituto Pirenaico de Ecologia (IPE-CSIC)
iRON	3	Osenga et al. (2019)
LAB-net	2	Mattar et al. (2016)
MySMNet	1	University Technology Malaysia
OZNET	13	Smith et al. (2012)
PBO_H2O	123	Larson et al. (2008)
REMEDHUS	5	http://campus.usal.es/~hidrus/
RSMN	19	http://assimo.meteoromania.ro/
SCAN	174	http://www.wcc.nrcs.usda.gov/
SMOSMANIA	19	Calvet et al. (2007), Albergel et al. (2008)
SNOTEL	256	http://www.wcc.nrcs.usda.gov/
SOILSCAPE	6	Moghaddam et al. (2010)

TERENO	3	Zacharias et al. (2011)
USCRN	99	Bell et al. (2013)

396

397 **2.3 Triple/Quadruple Collocation Analysis**

398 This section briefly introduces two Collocation-based Analyses, i.e., TCA and
 399 QCA, and the approach to estimate time-variant errors that have the same meaning of
 400 the TCA errors using ground measurements. TCA provides SM error estimates for
 401 three SM datasets without requiring an absolute truth. It assumes a linear relationship
 402 between SM observation θ_i and hypothetical unknown SM truth θ , which can be
 403 written as follows:

$$\theta_i = \alpha_i + \beta_i\theta + \varepsilon_i \quad i \in \{\text{GLDAS2, ASCAT, passive SM}\} \quad (1)$$

404 where the ε_i represents zero-mean random noise in θ_i , the α_i and β_i are additive and
 405 multiplicative coefficients that represent systematic errors in θ_i . The passive SM
 406 includes SMOSL3, SMOSIC, SMAPL3, and SMAPIB SM products. Based on this
 407 error model, SM errors can be obtained by calculating corresponding (co-)variances
 408 of the three SM datasets and simplifying with the other two assumptions, i.e., error
 409 orthogonality and zero Error Cross Correlation (ECC). Detailed formula derivation
 410 can be found in Gruber et al. (2016a).

411 Rescaling coefficients can be obtained through TCA along with error estimates.
 412 They can be used to make errors comparable after TCA or rescale SM data into a
 413 preselected reference dataset (Yilmaz and Crow, 2014; Gruber et al., 2017). Here,
 414 GLDAS2 was regarded as the reference dataset that was already perfectly calibrated.
 415 Therefore, the rescaling coefficients can be written as

$$\begin{cases} \beta_{\text{ASCAT}}^{\text{GLDAS2}} = \frac{\text{Cov}(SM_{\text{GLDAS2}}, SM_{\text{passive}})}{\text{Cov}(SM_{\text{ASCAT}}, SM_{\text{passive}})} \\ \beta_{\text{passive}}^{\text{GLDAS2}} = \frac{\text{Cov}(SM_{\text{GLDAS2}}, SM_{\text{ASCAT}})}{\text{Cov}(SM_{\text{passive}}, SM_{\text{ASCAT}})} \end{cases} \quad (2)$$

416 where $\beta_{\text{ASCAT}}^{\text{GLDAS2}}$ and $\beta_{\text{passive}}^{\text{GLDAS2}}$ are used to linearly rescale the ASCAT and passive-
 417 based SM errors against the GLDAS2 errors; $\text{Cov}(X, Y)$ represents covariance
 418 between X and Y time series.

419 The QCA extends TCA to four SM products and can account for the existence of
 420 non-zero ECC within a certain pair of collocated SM datasets. The QCA was used
 421 here to provide ECC values between GLDAS2, ASCAT, and SMOS/SMAP SM
 422 products and serve as an alternative error estimate method. The formulation reported
 423 in Gruber et al. (2016b) was conducted at the 759 sparse ground observation sites
 424 where ground measurements can serve as the fourth SM product. The least square
 425 solution for the QCA problem is given by

$$426 \quad \mathbf{y} = \begin{bmatrix} \sigma_a^2 \\ \sigma_b^2 \\ \sigma_c^2 \\ \sigma_d^2 \\ \sigma_{ab} \\ \frac{\sigma_{ac}\sigma_{ad}}{\sigma_{cd}} \\ \frac{\sigma_{bc}\sigma_{bd}}{\sigma_{cd}} \\ \frac{\sigma_{ad}}{\sigma_{ac}\sigma_{cd}} \\ \frac{\sigma_{bc}\sigma_{cd}}{\sigma_{bd}} \\ \frac{\sigma_{bd}}{\sigma_{ad}\sigma_{cd}} \\ \frac{\sigma_{ac}}{\sigma_{bd}\sigma_{cd}} \\ \frac{\sigma_{bc}}{\sigma_{ac}\sigma_{bd}} \\ \frac{\sigma_{cd}}{\sigma_{ad}\sigma_{bc}} \\ \sigma_{cd} \end{bmatrix} \quad \mathbf{A} = \begin{bmatrix} 1 & 0 & 0 & 0 & 0 & 1 & 0 & 0 & 0 & 0 \\ 0 & 1 & 0 & 0 & 0 & 0 & 1 & 0 & 0 & 0 \\ 0 & 0 & 1 & 0 & 0 & 0 & 0 & 1 & 0 & 0 \\ 0 & 0 & 0 & 1 & 0 & 0 & 0 & 0 & 1 & 0 \\ 0 & 0 & 0 & 0 & 1 & 0 & 0 & 0 & 0 & 1 \\ 1 & 0 & 0 & 0 & 0 & 0 & 0 & 0 & 0 & 0 \\ 0 & 1 & 0 & 0 & 0 & 0 & 0 & 0 & 0 & 0 \\ 0 & 0 & 1 & 0 & 0 & 0 & 0 & 0 & 0 & 0 \\ 0 & 0 & 1 & 0 & 0 & 0 & 0 & 0 & 0 & 0 \\ 0 & 0 & 0 & 1 & 0 & 0 & 0 & 0 & 0 & 0 \\ 0 & 0 & 0 & 1 & 0 & 0 & 0 & 0 & 0 & 0 \\ 0 & 0 & 0 & 0 & 1 & 0 & 0 & 0 & 0 & 0 \\ 0 & 0 & 0 & 0 & 1 & 0 & 0 & 0 & 0 & 0 \\ 0 & 0 & 0 & 0 & 1 & 0 & 0 & 0 & 0 & 0 \end{bmatrix} \quad \mathbf{x} = \begin{bmatrix} \beta_a^2 \sigma_\theta^2 \\ \beta_b^2 \sigma_\theta^2 \\ \beta_c^2 \sigma_\theta^2 \\ \beta_d^2 \sigma_\theta^2 \\ \beta_a \beta_b \sigma_\theta^2 \\ \sigma_{\varepsilon_a}^2 \\ \sigma_{\varepsilon_b}^2 \\ \sigma_{\varepsilon_c}^2 \\ \sigma_{\varepsilon_d}^2 \\ \sigma_{\varepsilon_a \varepsilon_b} \end{bmatrix} \quad (3)$$

427 where a, b, c, and d denote ASCAT, passive-based, GLDAS2, and ground SM,

428 respectively. To ensure the robustness of error estimates, TCA and QCA values were
 429 used only when Pearson's correlation coefficients between the triplet in TCA or
 430 quadruplet in QCA are greater than 0.2 and passed a t-test ($\alpha < 0.05$) (Scipal et al.,
 431 2008a, Su et al., 2014a). The least squares solution for the parameter \mathbf{x} is given as

$$\hat{\mathbf{x}} = (\mathbf{A}^T \mathbf{A})^{-1} \mathbf{A}^T \mathbf{y} \quad (4)$$

432 Ideally, ECC values between the SM triplet should be zero, which is a strong
 433 assumption in TCA method. In practice, researchers try to solve this problem by
 434 choosing three independent SM datasets, e.g., one active microwave, one passive
 435 microwave, and one model-based SM (Scipal et al., 2008a; Gruber et al., 2015;
 436 Gruber et al., 2016a; Wu et al., 2018). However, a certain level of ECC is inevitable
 437 in practice, especially for active and passive SM retrievals as reported in Gruber et al.
 438 (2016b) and Pierdicca et al. (2017). Therefore, it is necessary to inspect ECC values
 439 between these two kinds of SM datasets prior to TCA to guarantee reliable results.

440 As reported in Gruber et al. (2016b) and Chen et al. (2018), the ECC values can
 441 be obtained by the QCA. In particular, a selected data pair (e.g., ASCAT and passive-
 442 based SM here) is allowed to be correlated during the QCA while ECC between other
 443 data pairs are still required to be zero. Consequently, ECC value between the selected
 444 data pair can be obtained via their error covariance (i.e., $\sigma_{\varepsilon_a \varepsilon_b}$) and their error
 445 variance (i.e., $\sigma_{\varepsilon_a}^2$ and $\sigma_{\varepsilon_b}^2$) values derived from the Eq. (3). Alternatively, ECC values
 446 can be calculated based on the ground measurements directly. The boxplots of ECC
 447 values between ASCAT and passive SM datasets derived from QCA and ground
 448 measurements were shown in Fig. S2. For the QCA implementations consider

449 SMOSL3, SMOSIC, SMAPL3, and SMAPIB as the passive SM in the quadruplet, the
450 interquartile range is [-0.08, 0.34] and [-0.16, 0.35] for ECC values derived from
451 QCA and ground measurements, respectively, and the associated median values are
452 0.16 and 0.20 for these two cases. The small ECC values suggest that the strong
453 assumption in TCA is generally fulfilled in our analysis. It is notable that ECC values
454 can also be obtained by an extended double instrumental variable algorithm using
455 only two independent products, which was recently proposed by Dong et al. (2020).

456 TCA can provide time-variant errors by relaxing the stationary assumption. For
457 example, the time-variant errors can be obtained by applying TCA to SM time series
458 for a 30-day moving window advancing by 15-day steps over the experimental
459 periods (Loew and Schlenz, 2011). However, we argue that a moving-time-window
460 advancing by daily step can better capture the temporal variability of SM errors and
461 provide more inherent information about time-variant errors. Moreover, Su et al.
462 (2014) proposed an approach of estimating multi-annual window-based errors for
463 each day of the year. But this may overlook subtle temporal variability in SM errors
464 as the interannual variation of environmental conditions, such as vegetation and
465 rainfall, may change a lot in different years. Based on these considerations, a moving-
466 window-based TCA scheme was considered here.

467 Following the scheme proposed by Wu et al. (2021), we obtained ASCAT and
468 SMOS/SMAP time-variant daily errors by applying TCA to SM time series with a
469 sliding 101-day moving window advancing by daily step over the whole experimental
470 period. The 101-day time window was used here to estimate time-variant errors with

471 sufficient statistical power while keeping the kernel size of the window short enough
472 to capture the seasonal variability of SM errors. Since temporal interpolation was
473 considered for the SMOS and SMAP SM datasets here, the number of available triplet
474 samples used in the 101-day window is typically large enough to obtain reliable time-
475 variant errors. To avoid unreliable error estimates derived from limited samples (e.g.,
476 5 triplets), the minimum sample requirement was defined as 100 (Scipal et al., 2008a)
477 and 90 (Wu et al., 2021) for the time-invariant and time-variant error estimates,
478 respectively.

479 **2.4 Rescaling techniques and input selection in TCA**

480 Rescaling is an essential technique included in TCA as it eliminates bias between
481 different SM datasets prior to the TCA or adjusts the TCA error estimates into a
482 preselected reference space for a comparison purpose after the TCA implementation.
483 Here, we investigated four popular rescaling techniques generally applied in TCA:
484 three rescaling techniques applied prior to TCA, including variance and mean
485 matching (VAR), normalization (NORM), and CDF matching (CDF), and one
486 rescaling technique considers coefficients derived from TCA (TCA_Self). For
487 convenience, these abbreviations will be used to refer to the four rescaling techniques
488 described above. Since Section 2.3 has described the TCA_Self in equation (2), we
489 briefly introduced the other three rescaling techniques here.

490 VAR is a linear rescaling technique that forces SM time series to have the same
491 mean and standard deviation as the reference SM time series (Draper et al., 2009;
492 Brocca et al., 2010; Dorigo et al., 2010). Equation (5) describes this rescaling

493 approach.

$$SM_{\text{rescaled}} = \frac{SM - \mu(SM)}{\sigma(SM)} \sigma(SM_{\text{GLDAS2}}) + \mu(SM_{\text{GLDAS2}}) \quad (5)$$

494 where SM and SM_{rescaled} denote SM time series before and after rescaling,
495 respectively, $\mu(\cdot)$ and $\sigma(\cdot)$ represent mean and standard deviation.

496 NORM is a standardization method that forces SM time series to have the same
497 maximum and minimum as the reference SM time series (Rüdiger et al., 2009;
498 Albergel et al., 2010; Albergel et al., 2012; Su et al., 2014a; Wu et al., 2021). This
499 transformation can be written as the following equation (6).

$$SM_{\text{rescaled}} = \frac{SM - \min(SM)}{\max(SM) - \min(SM)} [\max(SM_{\text{GLDAS2}}) - \min(SM_{\text{GLDAS2}})] + \min(SM_{\text{GLDAS2}}) \quad (6)$$

500 where SM and SM_{rescaled} have the same meaning as those in equation (6), $\min(\cdot)$ and
501 $\max(\cdot)$ represent minimum and maximum values in corresponding time series,
502 respectively.

503 CDF can be considered as an enhanced nonlinear rescaling technique, which is
504 rescaled in such a way that the cumulative distribution function of SM time series is
505 matched with that of reference SM data (Reichle and Koster, 2004; Drusch et al. 2005;
506 Brocca et al. 2011; Brocca et al., 2013; Su et al., 2013). Equation (7) describes this
507 transformation.

$$CDF(SM_{\text{rescaled}}) = CDF(SM) \quad (7)$$

508 where $CDF(\cdot)$ represents cumulative distribution function, SM and SM_{rescaled} have
509 the same meaning as those in equation (5). Following Brocca et al. (2013), a fifth-

510 order polynomial function was considered in our CDF rescaling to match the two
 511 cumulative distribution functions.

512 Besides the aforementioned four rescaling techniques, QCA errors were also
 513 considered to make a comparison with the TCA errors. The ASCAT and
 514 SMOS/SMAP errors estimated by QCA were linearly rescaled against GLDAS2
 515 errors using the following equation

$$\left\{ \begin{array}{l} \sigma_{ASCAT_{after}} = \frac{\beta_{GLDAS2}}{\beta_{ASCAT}} \times \sigma_{ASCAT_{before}} \\ \sigma_{passive_{after}} = \frac{\beta_{GLDAS2}}{\beta_{passive}} \times \sigma_{passive_{before}} \quad passive \in \{SMOS, SMAP\} \end{array} \right. \quad (8)$$

516 where $\sigma_{ASCAT_{after}}$ and $\sigma_{passive_{after}}$ denote rescaled errors for ASCAT and passive SM
 517 data, i.e., SMOSL3, SMOSIC, SMAPL3, and SMAPIB SM products, $\sigma_{ASCAT_{before}}$ and
 518 $\sigma_{passive_{before}}$ represent ASCAT and passive SM errors before rescaling. The rescaling
 519 coefficients, $\frac{\beta_{GLDAS2}}{\beta_{ASCAT}}$ and $\frac{\beta_{GLDAS2}}{\beta_{passive}}$, can be obtained from equation (3).

520 TCA can be applied to original SM values or anomalies. Here, following Dorigo
 521 et al. (2010) and Albergel et al. (2012), anomalies were obtained by subtracting
 522 averages of SM values in a sliding time window from original SM values. The sliding
 523 window has a kernel size of 35-day and advances by daily step over the whole
 524 investigation period.

525 **2.5 Conventional ground-based approach to estimate SM errors**

526 To assess the evaluation power of TCA, TCA errors derived from the four
 527 rescaling techniques were validated by conventional SM errors obtained from ground
 528 measurements that are considered as the benchmark during the validation. As reported
 529 in Yilmaz and Crow (2014), ground-based error variance (regarded as the error term

530 in our work) can be obtained by equation (9), which has the same meaning as the
 531 TCA errors.

$$\sigma_{\varepsilon_X}^2 = \frac{1}{N} \sum_{t=1}^N (X_t - SM_{\text{Ground}2_t})^2 \quad X \in \{SM_{\text{SMOS}}, SM_{\text{SMAP}}, SM_{\text{ASCAT}}\} \quad (9)$$

532 where $\sigma_{\varepsilon_X}^2$ denotes ground-based error variance for the given dataset X , X_t and
 533 $SM_{\text{Ground}2_t}$ represent SM values of dataset X and ground measurement at time step t ,
 534 respectively, N is the number of whole investigation days.

535 In Eq. (9), the pre-processing of the ground measurements is different for each
 536 rescaling technique. For VAR, NORM, and CDF applied prior to TCA, ground
 537 measurements and other SM products were jointly rescaled by the same rescaling
 538 technique. By contrast, ground measurements were directly used without further
 539 process in TCA_Self and QCA_Self, and the resulting ground-based errors were
 540 rescaled by the same parameter and the same way as in the rescaling of TCA- and
 541 QCA-based errors.

542 **2.6 Robustness of temporal interpolation applied in TCA time-variant scheme**

543 Here, the TCA time-variant scheme was applied to passive-based SM datasets,
 544 i.e., SMOSL3, SMOSIC, SMAPL3, and SMAPIB, with temporal interpolation. The
 545 temporal interpolation introduces additional errors into SM time series and therefore
 546 may have potential impacts on the final TCA error estimates. Before applying TCA to
 547 temporally interpolated SM data, it is necessary to investigate such impact on the
 548 TCA time-variant errors.

549 To this end, first, we found a few pixels that have enough SM samples without

550 interpolation to calculate the TCA time-variant errors using a 101-day moving-time-
551 window. Since multiple passive SM datasets were considered in our work, the pixels
552 were separately selected for each TCA implementation that uses SMOSL3, SMOSIC,
553 SMAPL3, and SMAPIB SM as one of the triplet inputs. Second, the passive SM time
554 series were resampled without replacement using a different percentage of the original
555 data amount, i.e., 95%, 90%, 80%, 70%, 60%, and 50%. The passive SM time series
556 were then temporally interpolated and time-variant TCA errors were calculated based
557 on the resampled and interpolated SM datasets. The resulting time-variant errors were
558 compared with the time-variant errors derived from the original SM data (without
559 interpolation) using the Pearson's correlation and RMSD metrics. To guarantee
560 reliable conclusions, the resampling was repeated 1000 times and we took an average
561 of the resulting correlation coefficients and RMSD values to represent the overall
562 performance of interpolation on TCA error estimates. Finally, correlations and RMSD
563 values between time-variant errors calculated from original SM data and SM time
564 series resampled by different percentage values can be obtained for each selected
565 pixel. Based on the correlation and RMSD results, the impact of temporal
566 interpolation on TCA errors can be quantified.

567 **2.7 Evaluation metrics**

568 The Relative Difference (RD) and Overall Relative Difference (ORD) metrics
569 were used here to evaluate the relative difference between time-invariant and time-
570 variant errors. In a given pixel, the ORD is defined as:

$$\text{ORD} = \sqrt{\frac{1}{N} \sum_{t=1}^N \left(\frac{TVE_t - TIE}{TIE} \right)^2} \times 100\% \quad (10)$$

571 where TIE denotes time-invariant error and TVE_t represents time-variant error at time
 572 step t , N denotes the actual number of available common pairs of TIE and TVE in the
 573 given pixel. ORD quantifies the overall relative difference between time-invariant and
 574 time-variant errors. But it fails to provide intuitive information about the positive or
 575 negative sign of the relative difference between these two kinds of errors. Therefore,
 576 the RD metric was constructed to address this issue. For a given pixel, the RD is
 577 defined as follows:

$$\text{RD} = \frac{1}{N} \sum_{t=1}^N \frac{TVE_t - TIE}{TIE} \times 100\% \quad (11)$$

578 where symbols in equation (11) are the same as those in equation (10).

579 Pearson's correlation and RMSE metrics were used to evaluate the accuracy of
 580 TCA-based error estimates and ground-based errors were regarded as the benchmark.
 581 The RMSE is defined as follows:

$$\text{RMSE} = \sqrt{\frac{1}{N} \sum_{t=1}^N (\sigma_{\text{TCA}_t} - \sigma_{\text{ground}_t})^2} \quad (12)$$

582 where σ_{ground} is ground-based errors and σ_{TCA} represents TCA-based errors, N
 583 denotes the number of associated common pairs of σ_{ground} and σ_{TCA} during the
 584 investigation period.

585 The calculation of RMSD metric is similar to RMSE and is defined as:

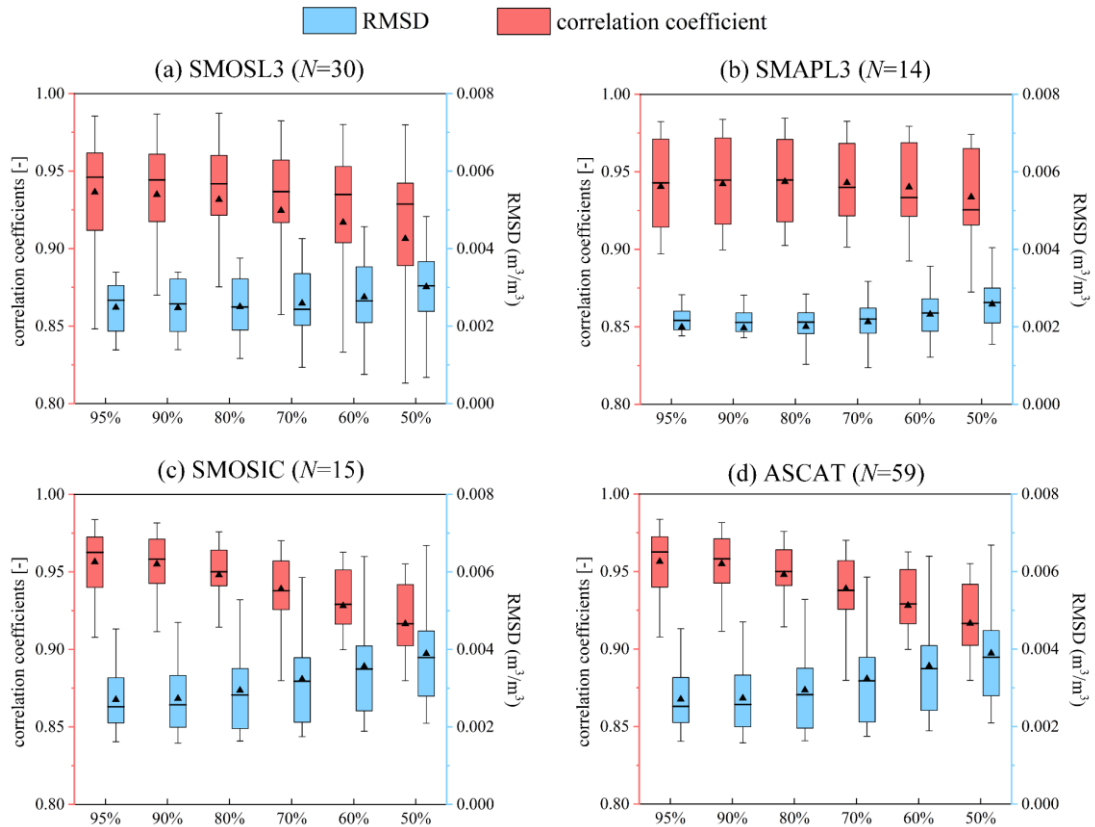
$$\text{RMSD} = \sqrt{\frac{1}{N} \sum_{t=1}^N (\sigma_{\text{resampled}_t} - \sigma_{\text{original}_t})^2} \quad (13)$$

586 where $\sigma_{\text{resampled}}$ denotes TCA time-variant errors derived from SM time series
 587 resampled by the different percentage values whereas σ_{original} represents TCA time-
 588 variant errors estimated from original SM data without interpolation.

589 **3 Results**

590 **3.1 Impact of temporal interpolation on TCA error estimates**

591 As noted in Sec. 2.6, a potential error source of the TCA estimates in our work is
 592 the temporal interpolation applied to the passive SM datasets, which may lead to a
 593 potential impact on the final TCA error estimates. Fig. 2 explores such impact using
 594 the method described in Sec. 2.6. SMAPIB results are not included here as limited
 595 available samples (< 30) were found. Three ASCAT results are expected from the
 596 TCA implementations with SMOSL3, SMAPL3, and SMOSIC, and all of them are
 597 used to construct the boxplots in Fig. 2 (d).



598

599 Fig. 2 Boxplots of correlation coefficients (red) and RMSD values (blue) between
600 time-variant errors derived from the original SM data without interpolation and the
601 SM time series resampled by different percentage values of the original data. (a-d)
602 exhibits the results of applying interpolation to SMOSL3, SMAPL3, SMOSIC, and
603 ASCAT SM data, respectively. The x-axis denotes the percentage values of the
604 original data considered in the resampling. The y-axes on the left (red color) and right
605 (blue color) describe correlation coefficients and RMSD, respectively. The ‘*N*’ above
606 each subfigure denotes the number of available samples included in the corresponding
607 boxplot.

608 Boxplots in Fig. 2 demonstrate that the temporal interpolation has a small impact
609 on the final TCA error estimates. In line with expectations, for all cases shown in Fig.
610 2 (a-d), correlation displays a decreasing trend while RMSD yields an increasing

611 trend as the sampling percentage of the original dataset decreases, which indicates the
612 temporal interpolation indeed introduces additional errors into SM time series and
613 consequently degenerates the accuracy of the final TCA error estimates. However, it
614 is worthwhile mentioning that this degeneration is relatively small as correlation
615 coefficients were typically high (> 0.9) and RMSD values were generally small ($<$
616 $0.004 \text{ m}^3/\text{m}^3$) for all the cases. Moreover, as the sampling percentage decreases from
617 95% to 50%, the correlation decreased by a value smaller than 0.04 and the RMSD
618 typically increased by a value smaller than $0.001 \text{ m}^3/\text{m}^3$ regarding their average and
619 median values. These results demonstrate that the temporal interpolation applied in
620 our work has a small impact on TCA error estimates and it is an efficient way to
621 address the limited sample issue lies in the TCA time-variant scheme.

622 **3.2 Accuracy of TCA errors derived from different rescaling techniques**

623 Pearson's correlation coefficients and RMSE values between ground- and TCA-
624 based time-invariant errors for ASCAT and multiple passive SM products are
625 summarized in Table 3 and Table 4, respectively. The impacts of the four rescaling
626 techniques (i.e., VAR, NORM, CDF, and TCA_Self) and the two input selections (i.e.,
627 original values or anomalies) on the final TCA time-invariant errors are compared in
628 these two tables. Moreover, QCA error estimates (marked as QCA_Self) were also
629 considered to make a comparison with the TCA errors. The correlation and RMSE
630 values of ASCAT are averaged from multiple ASCAT results that are obtained from
631 the TCA implementations with multiple passive SM data used here.

632 Table 3 Pearson's correlation coefficients between ground- and TCA-based time-

633 invariant errors for SMOSL3, SMOSIC, SMAPL3, SMAPIB, and ASCAT SM
634 products. The time-invariant errors are derived from multiple rescaling techniques.
635 The first and second recommendation strategies (excluding QCA_Self) to implement
636 TCA are highlighted with green and yellow colors, respectively. All the ρ values
637 passed a t-test ($\alpha < 0.05$)

		NORM	VAR	CDF	TCA_Self	QCA_Self
original values	SMOSL3	0.25	0.39	0.42	0.65	0.92
	SMOSIC	0.15	0.31	0.40	0.54	0.97
	SMAPL3	0.19	0.20	0.34	0.38	0.96
	SMAPIB	0.21	0.19	0.16	0.50	0.74
	ASCAT	0.49	0.70	0.73	0.86	0.89
anomalies	SMOSL3	0.69	0.75	0.74	0.95	0.96
	SMOSIC	0.64	0.66	0.63	0.89	0.90
	SMAPL3	0.56	0.58	0.60	0.70	0.79
	SMAPIB	0.47	0.58	0.54	0.70	0.67
	ASCAT	0.74	0.79	0.79	0.93	0.90

638

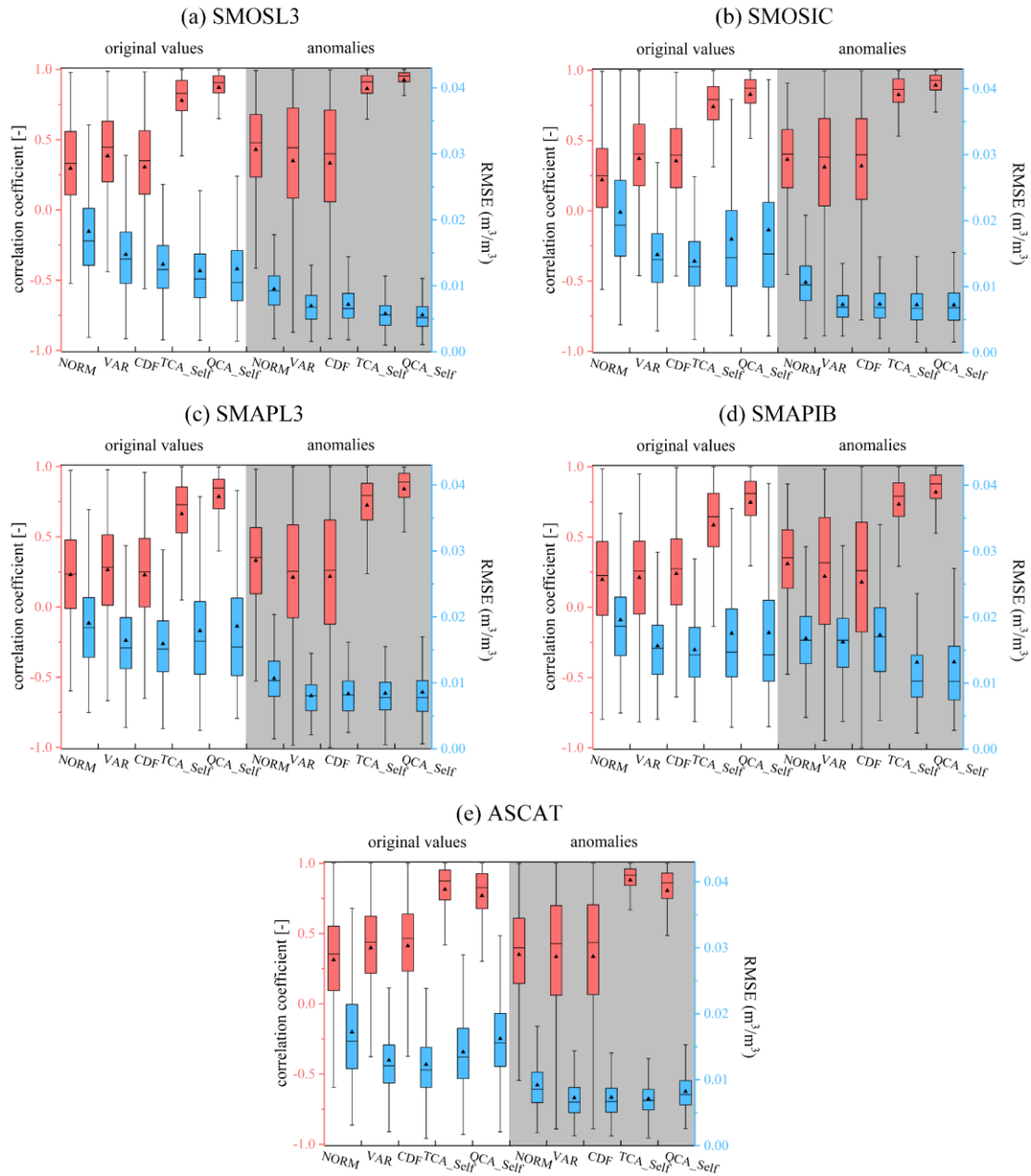
639 Table 4. The same as Table 3 but for the RMSE values

		NORM	VAR	CDF	TCA_Self	QCA_Self
original values	SMOSL3	0.0256	0.0191	0.0176	0.0178	0.0220
	SMOSIC	0.0295	0.0188	0.0178	0.0276	0.2208
	SMAPL3	0.0265	0.0227	0.0214	0.0263	0.0575
	SMAPIB	0.0265	0.0207	0.0203	0.0241	0.0266
	ASCAT	0.0231	0.0156	0.0147	0.0194	0.0217
anomalies	SMOSL3	0.0086	0.0072	0.0074	0.0047	0.0046
	SMOSIC	0.0099	0.0075	0.0077	0.0062	0.0062
	SMAPL3	0.0102	0.0085	0.0087	0.0083	0.0086
	SMAPIB	0.0163	0.0184	0.0214	0.0158	0.0157
	ASCAT	0.0079	0.0075	0.0077	0.0060	0.0070

640 Tables 3 and 4 jointly show that the highest correlation and the smallest RMSE
641 values between ground- and TCA-based time-invariant errors are obtained by
642 applying TCA to SM anomalies and rescaling the resulting errors using TCA_Self.

643 Table 3 shows the correlations between these two kinds of errors were 0.95, 0.89,
644 0.70, 0.70, and 0.93 for the SMOSL3, SMOSIC, SMAPL3, SMAPIB, and ASCAT
645 SM products, respectively. This combination strategy of applying anomalies and
646 TCA_Self was also recommended based on the RMSE results shown in Table 4, as
647 the smallest RMSE value was typically observed when considering this combination
648 strategy. Furthermore, Tables 3 and 4 indicate that there is a small difference between
649 errors derived from TCA_Self and QCA_Self when considering anomalies as the SM
650 inputs. By contrast, errors estimated from TCA_Self and QCA_Self exhibited evident
651 discrepancies for the correlation and RMSE metrics when considering absolute values
652 as inputs. In particular, as shown in Table 3, errors derived from QCA_Self appeared
653 higher correlation coefficient than those derived from TCA_Self when using absolutes
654 as SM inputs.

655 The validation of TCA time-variant errors derived from different rescaling
656 techniques using ground-based errors are revealed in Fig. 3. Specifically, TCA- and
657 ground-based time-variant errors are compared using Pearson's correlation and
658 RMSE metrics for each experimental grid cell. The correlation coefficients and
659 RMSE values were gleaned and construct the boxplots shown in Fig. 3. Since
660 correlation and RMSE values of ASCAT SM product can be obtained for each TCA
661 implementation that applied multiple passive SM datasets, all the ASCAT results are
662 considered in Fig. 3 (e).



663

664 Fig. 3 Boxplots of the Pearson's correlation coefficients (red) and RMSE values (blue)

665 between ground- and TCA-based time-variant errors for the TCA implementations

666 that consider (a) SMOSL3, (b) SMOSIC, (c) SMAPL3, (d) SMAPIB, and (e) ASCAT

667 as the triplet inputs. In each subfigure, results derived from absolutes and anomalies

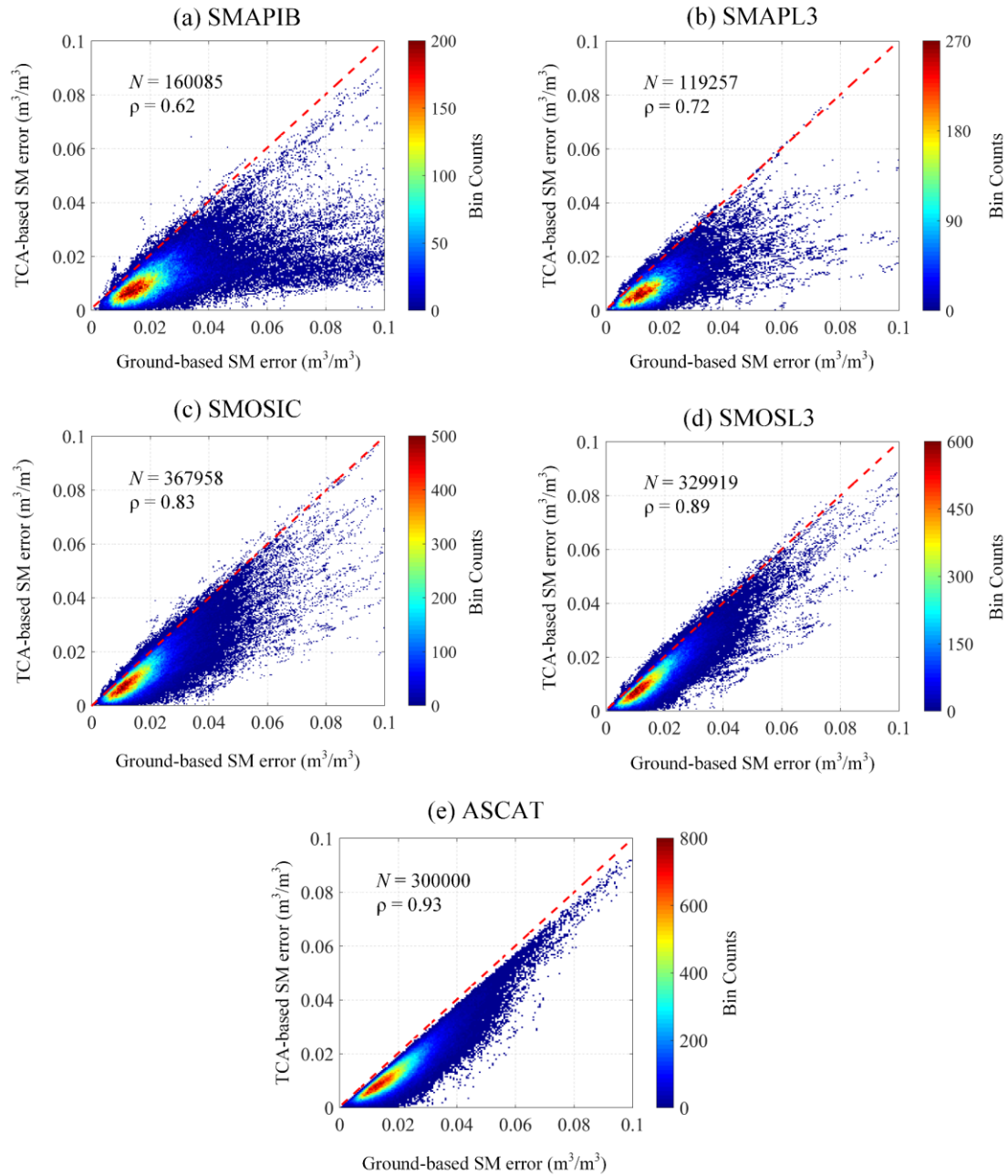
668 are shown in the white and grey areas, respectively. The x-axis is the multiple

669 rescaling techniques considered in the TCA. The y-axes on the left (red color) and

670 right (blue color) describe correlation coefficients and RMSE, respectively.

671 Similar to the conclusion of time-invariant errors drawn from Tables 3 and 4,
672 boxplots in Fig. 3 also demonstrate that the highest correlation and the smallest
673 RMSE values between ground- and TCA-based time-variant errors are typically
674 obtained by the combination strategy that considers SM anomalies as TCA inputs and
675 rescaling the resulting errors with TCA_Self. Based on this optimal combination
676 strategy, TCA errors were strongly correlated with ground-based errors as associated
677 correlation coefficients were 0.88 (0.92), 0.83 (0.87), 0.86 (0.91), 0.73 (0.79), and
678 0.72 (0.79) for the ASCAT, SMOSIC, SMOSL3, SMAPIB, and SMAPL3 cases
679 regarding the average (median) values. Corresponding scatterplots are illustrated in
680 Fig. 4 and all available samples were included in such comparison.

681 In contrast with the correlation metric, RMSE values derived from TCA_Self
682 were significantly smaller than those obtained from the other three rescaling
683 techniques (excluding QCA_Self) only for the SMOSL3 and SMAPIB cases.
684 However, it has to be stressed that TCA_Self RMSE values were generally small for
685 all cases, i.e., typically smaller than $0.015 \text{ m}^3/\text{m}^3$ for SMAPIB and smaller than 0.01
686 m^3/m^3 for SMOSL3, SMOSIC, and SMAPL3. Furthermore, Fig. 3 implies that
687 ground-based errors had a stronger correlation with QCA_Self errors than TCA_Self
688 errors for all cases except for the ASCAT case. However, one can see that the
689 discrepancies were relatively small for the TCA_Self and QCA_Self RMSE values.



690

691 Fig. 4 Scatterplots of ground- and TCA-based time-variant errors for (a) SMAPIB, (b)

692 SMAPL3, (c) SMOSIC, (d) SMOSL3, and (e) ASCAT SM data. The original scatter

693 points are binned in the increase of x-axis for ground-based error and y-axis for TCA-

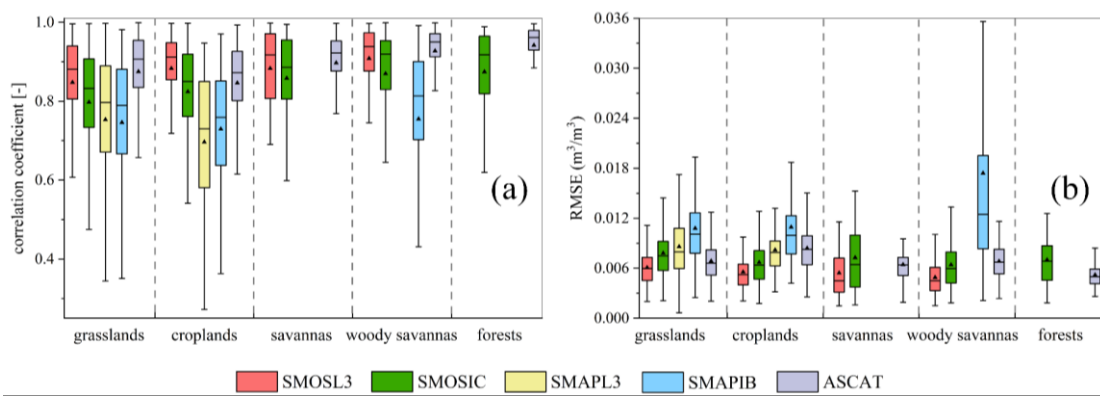
694 based error, respectively, and colored based on the number of points included in the

695 bin. All the ρ values passed a t-test ($\alpha < 0.05$). N describes the number of

696 experimental grid cells considered in each scatterplot.

697 The scatterplots in Fig. 4 exhibit that TCA-based errors are strongly correlated
 698 with ground-based errors as their ρ values were in the range from 0.62 to 0.93.
 699 However, it is notable that TCA- and ground-based errors had a relatively weak
 700 correlation ($\rho = 0.62$) for the SMAPIB case, and considerable points with an
 701 underestimation of the TCA-based errors were observed in Fig. 4 (a). Moreover, Fig.
 702 4 demonstrates that ground-based errors were generally greater than TCA-based
 703 errors. Given the above results, only TCA errors obtained from the optimal
 704 combination strategy, i.e., applying TCA to SM anomalies with TCA_Self, were
 705 considered in the following Sections 3.4 and 3.5.

706 At the end of this section, the correlation and RMSE values between TCA- and
 707 ground-based time-variant errors are further investigated in Fig. 5 by a classification
 708 with six land cover types, which explores the TCA performance in different land
 709 covers. Since multiple passive SM data were applied in TCA, all the resulting
 710 correlation and RMSE values of ASCAT SM data were considered in Fig. 5.



712 Fig. 5 Boxplots of (a) correlation coefficients and (b) RMSE between TCA- and
 713 ground-based errors classified by six land cover types for ASCAT and multiple
 714 passive SM products. Boxplots with limited available samples (< 30) are not

715 considered in such comparisons.

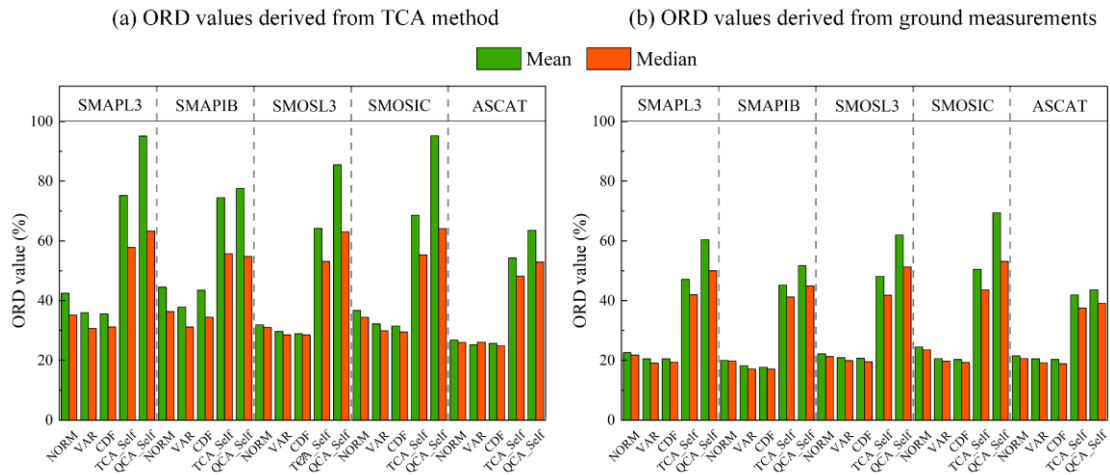
716 In general, Fig. 5 reveals that TCA can accurately estimate ASCAT and
717 SMOS/SMAP time-variant errors in all the six land cover types as associated
718 correlations are typically greater than 0.6 and the RMSE values are mostly smaller
719 than $0.01 \text{ m}^3/\text{m}^3$. Nevertheless, Fig. 5 (a) shows relatively weak correlations between
720 TCA- and ground-based errors in croplands for all SM datasets except for SMOSL3
721 SM, which implies TCA method has relatively less power to efficiently characterize
722 errors of satellite-based SM products in croplands.

723 In comparison, TCA provides more accurate error estimates for SMOS SM than
724 SMAP SM when considering ground-based errors as the benchmark. Even though
725 several boxplots were not included in Fig. 5 due to their limited sample issue, it
726 appears that the correlations presented a decreasing trend and the RMSE gave an
727 increasing trend for SM products in the order of SMOSL3, SMOSIC, SMAPL3, and
728 SMAPIB. Notably, the discrepancy between TCA- and ground-based errors is evident
729 for SMAPIB SM, especially in woody savannas.

730 **3.3 Relative difference between time-variant and time-invariant errors**

731 This section identifies the relative difference between time-variant and time-
732 invariant SM errors derived from different rescaling techniques to find out (i) the
733 necessity to consider and include time-variant SM errors in applications and (ii) which
734 one is larger, time-variant or time-invariant error? To these ends, the ORD and RD
735 metrics were used to address these two issues. For a given combination of passive SM
736 dataset and rescaling technique, the ORD value can be calculated for each

737 experimental grid cell. Mean and median of the ORD values collected from all
 738 experimental pixels were used to describe the overall deviation between time-
 739 invariant and time-variant errors. Note that only TCA errors obtained from SM
 740 anomalies were considered in this section.



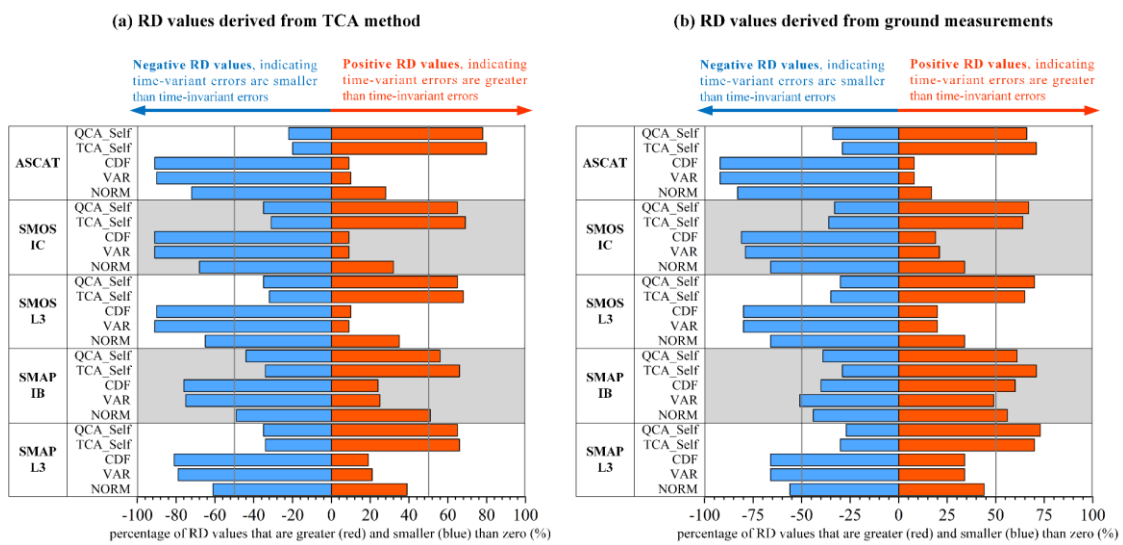
741

742 Fig. 6 Mean and median of the ORD values derived from the NORM, VAR, CDF,
 743 TCA_Self, and QCA_Self rescaling techniques for SMAPL3, SMAPIB, SMOSL3,
 744 SMOSIC, and ASCAT SM datasets. (a) and (b) represent ORD values obtained from
 745 the TCA- and ground-based evaluations, respectively.

746 The bar charts in Fig. 6 demonstrate that it is necessary to consider and include
 747 time-variant errors in actual applications. The ORD average and median values
 748 derived from different rescaling techniques were typically greater than 25% and 20%
 749 for the TCA-based and ground-based errors, respectively. In particular, ORD average
 750 and median values derived from the TCA_Self and QCA_Self methods were mostly
 751 greater than 50% and 40% for errors obtained from TCA and ground measurements,
 752 respectively. Comparing Fig. 6 (a) and (b), one can see that TCA-based ORD values
 753 were typically smaller than the ground-based ORD values, which implies TCA-based

754 time-variant errors have a larger variance than ground-based time-variant errors.

755 The first question proposed at the beginning of this section is explored using the
 756 ORD metric in the above analysis. However, the ORD metric only explores overall
 757 magnitude of the relative differences between time-variant and time-invariant errors.
 758 The RD metric is adopted here to answer the second question put forward in this
 759 section.



760

761 Fig. 7 Fractions of the RD values that are greater (red) or smaller (blue) than 0 for the
 762 combination of five SM datasets, i.e., ASCAT, SMOSIC, SMOSL3, SMAPIB, and
 763 SMAPL3, and five rescaling techniques, i.e., QCA_Self, TCA_Self, CDF, VAR, and
 764 NORM. (a) and (b) present associated results based on the RD values derived from
 765 the TCA- and ground-based evaluations, respectively.

766 The RD results in Fig. 7 reveal that the relative magnitude relationship between
 767 time-invariant and time-variant errors is varied with the selection of rescaling
 768 technique used in TCA. In general, negative RD values were observed for the CDF,
 769 VAR, and NORM rescaling techniques, implying time-variant errors were smaller

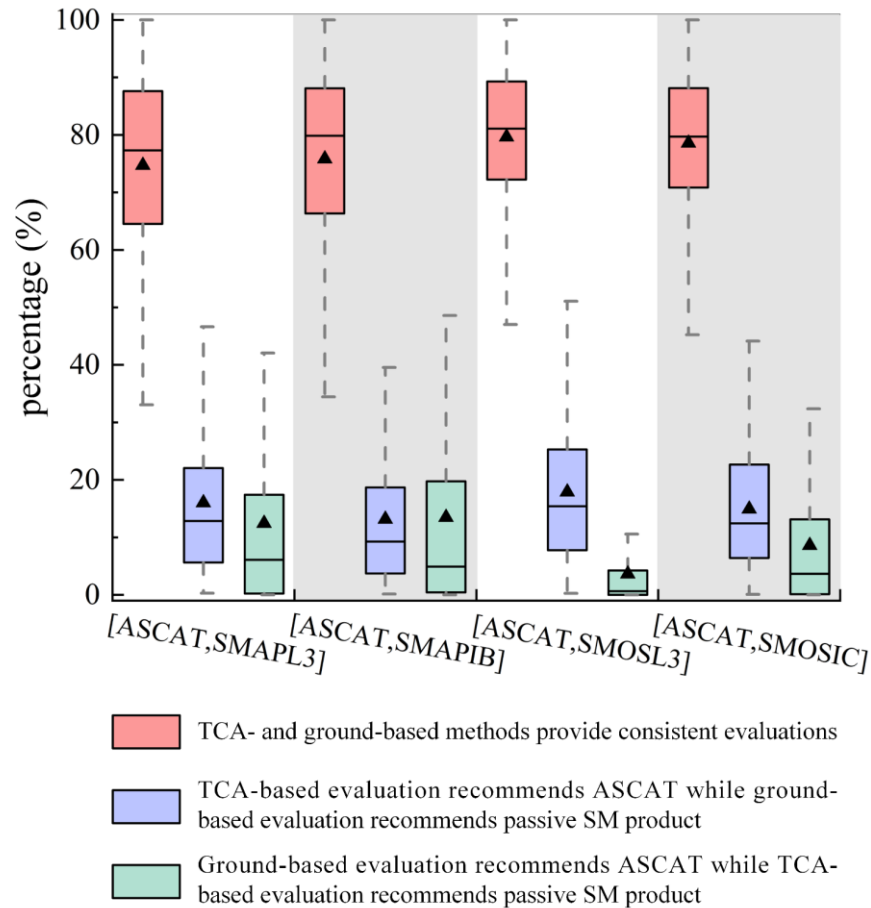
770 than time-invariant errors. By contrast, positive RD values were found for the
771 TCA_Self and QCA_Self rescaling techniques, which indicates time-variant errors
772 were greater than time-invariant errors. It is worth mentioning that the RD results of
773 SMAPIB SM are different from other cases, especially for the ground-based RD
774 values shown in Fig. 7 (b).

775 The noticeable difference between time-variant and time-invariant errors
776 demonstrated in Fig. 6 suggests that considering time-variant errors rather than time-
777 invariant errors in applications is necessary. Time-variant error better characterizes
778 the temporal variability of SM errors and consequently, a more accurate output is
779 expected from applications that strongly rely on an accurate error specification, such
780 as data merging and data assimilation studies. Contrary to the expectation, time-
781 invariant errors do not provide an average reference for time-variant errors derived
782 from TCA and their relative magnitude depends on the rescaling technique used.
783 Simple rescaling techniques, such as VAR, CDF, and NORM, tend to underestimate
784 SM time-variant errors when window-based TCA was applied.

785 **3.4 Evaluation consistency between TCA- and ground-based methods**

786 This section assesses the evaluation consistency between TCA- and ground-based
787 methods regarding the time-variant errors of ASCAT and multiple passive SM data
788 used in our work. For each experimental grid cell, the evaluation consistency was
789 quantified as a percentage value, which represents the fraction of days that TCA- and
790 ground-based methods provide consistent evaluation results. Four assessment results
791 can be obtained for the four TCA implementations that consider SMOSL3, SMOSIC,

792 SMAPL3, and SMAPIB as the passive SM in the triplet. In each TCA implementation,
 793 the consistent evaluation includes two cases: both TCA and ground-based methods
 794 suggest (i) ASCAT or (ii) the given passive SM dataset has the smallest time-variant
 795 error.



796

797 Fig. 8 Boxplots of a binary assessment of the evaluation consistency between TCA-
 798 and ground-based methods for ASCAT and four passive-based SM time-variant errors.
 799 The four columns illustrated in this figure denote the four TCA implementations that
 800 consider [GLDAS2, ASCAT, SMAPL3], [GLDAS2, ASCAT, SMAPIB], [GLDAS2,
 801 ASCAT, SMOSL3], and [GLDAS2, ASCAT, SMOSIC] as the triplet.

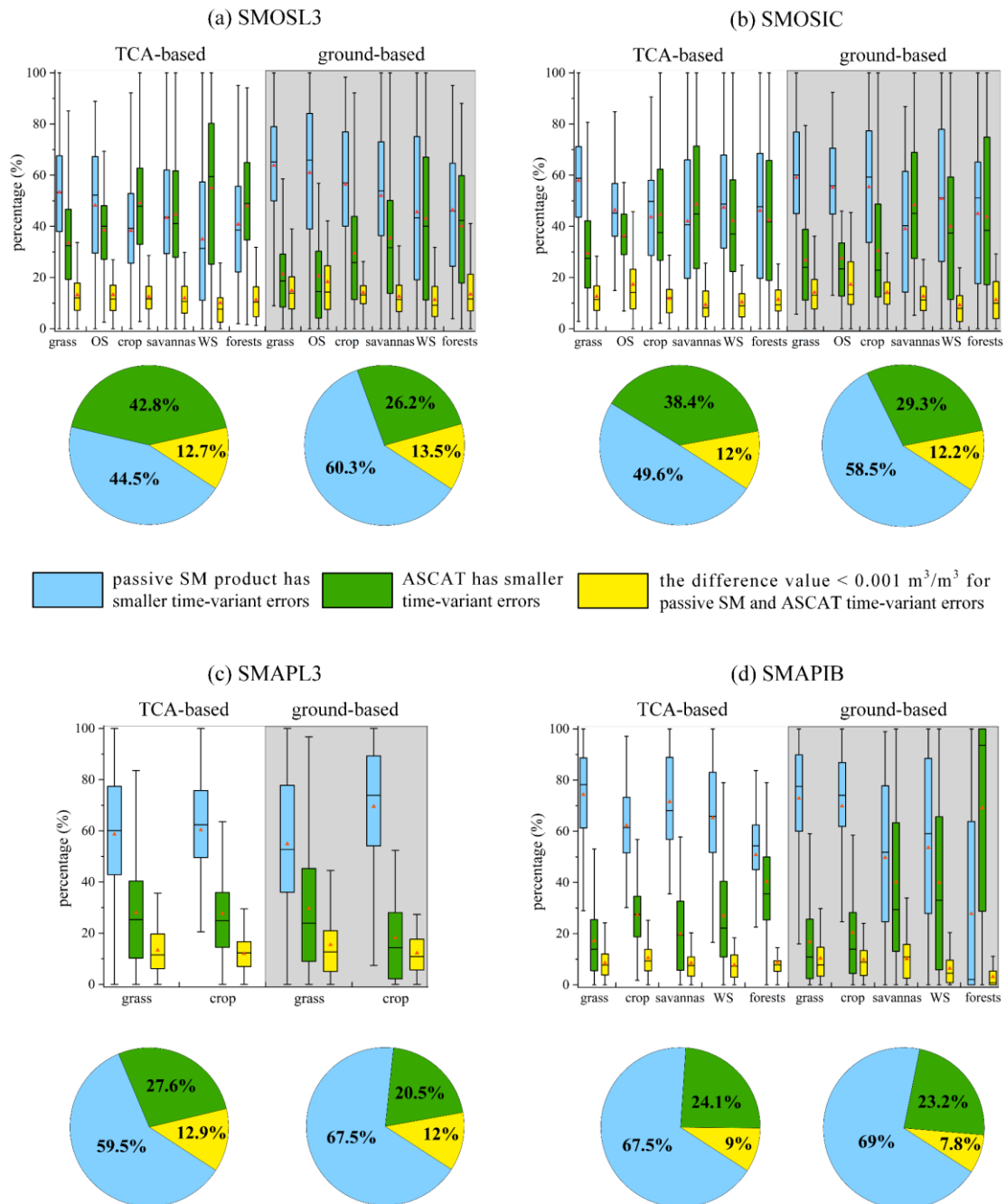
802 Boxplots in Fig. 8 show that there is a high consistency between TCA- and
 803 ground-based evaluations. These two kinds of methods provided consistent

804 evaluations in 74.7% (77.3%), 75.8% (79.8%), 79.6% (81.1%), and 78.6% (79.7%) of
805 the investigation days on the global average (median) for the TCA implementations
806 with SMAPL3, SMAPIB, SMOSL3, and SMOSIC SM, respectively. However, the
807 boxplots with dark blue and cyan in Fig. 8 also exhibit that the TCA- and ground-
808 based evaluations provided inconsistent evaluation results in certain periods. Notably,
809 TCA more or less underestimated ASCAT errors and overestimated passive-based
810 SM errors and this phenomenon was more pronounced for the SMOS data than the
811 SMAP data.

812 **3.5 Comparisons of time-variant errors by land cover classification for ASCAT** 813 **and four passive SM products**

814 This section is an extensive analysis of the evaluation consistency shown in Sec.
815 3.4 by pairwise comparing TCA- and ground-based time-variant errors for ASCAT
816 and the four passive SM products, i.e., SMOSL3, SMOSIC, SMAPL3, and SMAPIB.
817 However, this section focuses on the error comparison while Sec. 3.4 aims at
818 quantifying the evaluation consistency between TCA- and ground-based methods.
819 Also, this section provides insights into the relative strengths of ASCAT and the four
820 passive SM products in different land cover types regarding their time-variant errors.
821 Four comparison results are expected as TCA was applied to the four above passive
822 SM products. In such comparisons, three cases were considered: (i) ASCAT or (ii)
823 passive-based SM has the smallest time-variant error, and (iii) ASCAT has similar
824 performance with passive SM as the difference between their errors is smaller than
825 $0.001 \text{ m}^3/\text{m}^3$. For a given TCA implementation that uses SMOS or SMAP SM as the

826 triplet, three percentage values corresponding to the above three cases can be
 827 calculated for each grid cell. The percentage values were collected from all pixels and
 828 were further categorized by six land cover types including grasslands (grass), open
 829 shrublands (OS), croplands (crop), savannas, woody savannas (WS), and forests.



830

831 Fig. 9 Boxplots of the percentage values classified by six land cover types for the

832 three cases: ASCAT (green boxes) or passive SM (blue boxes) has the smallest time-

833 variant error and the difference between their errors is smaller than $0.001 \text{ m}^3/\text{m}^3$
834 (yellow boxes). (a-d) represent the TCA- (white areas) and ground-based (grey areas)
835 assessments that consider SMOSL3, SMOSIC, SMAPL3, and SMAPIB SM data,
836 respectively. To guarantee reliable results, boxplots associated with limited available
837 samples (< 30) are excluded. The two pies below each subfigure exhibit the overall
838 performance of the three cases that include all available samples for the TCA- (on the
839 left) and ground-based (on the right) assessments.

840 Overall, both TCA- and ground-based comparisons demonstrate that ASCAT
841 provides more and more reliable SM observations as vegetation cover increases while
842 passive-based SM provides more and more reliable observations as vegetation cover
843 decreases. The boxplots in Fig. 9 exhibit that the percentage value typically becomes
844 smaller for passive SM while getting larger for ASCAT SM as vegetation cover
845 increases. For all the four cases illustrated in Fig. 9 (a-d), percentage values of the
846 case that ASCAT provided similar performance with passive SM (the yellow boxes)
847 had relatively small differences in the six land cover types, and associated percentage
848 values were mostly smaller than 20%.

849 The pie charts in Fig. 9 confirm that TCA underestimated ASCAT errors and
850 overestimated SMOS errors when considering ground-based evaluation as the
851 benchmark. Comparing the pie charts on the left and right sides, the fraction value
852 that ASCAT had better performance decreased by 16.6%, 9.1%, 7.1%, and 0.9% for
853 the SMOSL3, SMOSIC, SMAPL3, and SMAPIB cases, respectively. Conversely,
854 fraction value that passive-based SM provided smaller time-variant error increased by

855 15.8%, 8.9%, 8%, and 1.5% for the above four cases, respectively. However, this
856 phenomenon is not evident for SMAPIB when compared with the SMOSL3,
857 SMOSIC, and SMAPL3.

858 Based on the land cover analysis in Fig. 9, we can further investigate the land
859 cover types where TCA was found to underestimate ASCAT errors whereas
860 overestimate passive SM errors. As shown in Fig. 9 (a), this phenomenon was
861 observed in all the six land cover types for the SMOSL3 SM. However, this mismatch
862 phenomenon was only evident in open shrublands and croplands for the SMOSIC SM
863 illustrated in Fig. 9 (b). Although the available samples are relatively few for
864 SMAPL3 in Fig. 9 (c), one can still see that the mismatch phenomenon was obvious
865 in croplands. Compared with the above three SM products, the TCA performance is
866 more complex for the SMAPIB SM. Fig. 9 (d) shows this mismatch phenomenon was
867 prominent in croplands for the SMAPIB case. However, TCA was found to
868 overestimate ASCAT errors and underestimate SMOS errors for the SMAPIB in
869 savannas, woody savannas, and forests when considering ground-based errors as the
870 benchmark.

871 **4 Discussion**

872 It is necessary to consider the rescaling technique used in TCA as the selection of
873 rescaling method has a great impact on the accuracy of the final TCA error estimates,
874 which is generally ignored in current TCA studies. Ground- and TCA-based errors
875 achieve the highest correlation and the smallest RMSE when TCA is applied to SM
876 anomalies with the TCA_Self rescaling technique. Rescaling the inputs against a

877 reference prior to TCA can implant the distribution information of the reference
878 dataset to all TCA inputs. Consequently, the input SM products to TCA may show
879 spurious large cross-correlation and make the resulting TCA errors smaller as well.
880 By contrast, TCA_Self does not affect the original SM observations during the TCA
881 calculation. It provides optimal rescaling parameters as it considers the individual
882 random error properties and matches the variability of the jointly observed signal
883 (Gruber et al., 2017). Although TCA_Self and other suboptimal rescaling techniques
884 can address the first-order (additive) biases for considered datasets. However, the
885 TCA_Self can address the second-order (multiplicative) biases while other suboptimal
886 rescaling techniques cannot (Gruber et al., 2020). This can explain the better
887 performance of TCA_Self than other rescaling techniques.

888 TCA_Self was also recommended to use in data assimilation to remove the
889 systematic bias between model simulation and observations when compared with
890 VAR and CDF rescaling techniques (Yilmaz and Crow, 2013). However, TCA_Self
891 was applied prior to TCA to rescale SM inputs in Yilmaz and Crow (2013) and error
892 estimates were derived from the rescaled SM time-series. By contrast, the optimal
893 strategy proposed in our work applied TCA_Self after the TCA implementation to
894 rescale the error estimates. These two kinds of approaches are supposed to lead to the
895 same error estimates (Gruber et al., 2016a). Our result is complementary to the study
896 by Yilmaz and Crow (2013) and a synergy result can be drawn that TCA_Self is an
897 optimal rescaling technique not only for removing systematic bias between SM
898 datasets but also for estimating errors using the TCA method.

899 The ORD results suggest considering and including time-variant SM errors in
900 applications as time-variant errors deviate from time-invariant errors by over 50%
901 when TCA_Self is used. Distinct temporal variability of SM errors at short time scales
902 such as seasonally, monthly, or daily time scales was confirmed in Loew and Schlenz
903 (2011), Zwieback et al. (2012), Su et al. (2014a), and Wu et al. (2021). However, we
904 found the relative magnitude relationship between time-invariant and time-variant
905 errors is not the same for different rescaling techniques used in TCA, which has not
906 been fully addressed by previous studies. Loew and Schlenz (2011) and Wu et al.
907 (2021) found time-invariant errors are greater than time-variant errors. Our results
908 confirmed this conclusion when SM products are rescaled against a reference dataset
909 prior to TCA. However, time-invariant errors are smaller than time-variant errors
910 when error estimates are rescaled by TCA_Self. This result indicates that simple
911 matching technique, including NORM, VAR, and CDF, tends to underestimate time-
912 variant errors when less and less data pairs are considered in the moving-window-
913 based TCA. Moreover, SM inputs are rescaled against a reference dataset in the
914 simple matching techniques, which can significantly influence the magnitude of the
915 final TCA errors (Draper et al., 2013; Dorigo et al., 2015).

916 Even though we found a strong consistency between TCA- and ground-based
917 evaluations, these two evaluations have distinct differences in two aspects. First,
918 ground-based errors are larger than TCA-based errors for both time-invariant and
919 time-variant cases, which is in line with Yilmaz and Crow (2014) and Dorigo et al.
920 (2015). Second, TCA typically overestimates passive-based SM errors and

921 underestimates ASCAT errors when considering ground-based errors as the
922 benchmark. The better TCA results for ASCAT SM could be due to its relatively
923 better match with GLDAS2 SM as they have more consistent spatial
924 representativeness in grid cells. The spatial representativeness of ASCAT and passive
925 SM datasets is kept since we used the nearest neighbor resampling. This resampling
926 method can preserve original SM values in an unaltered way but may introduce the
927 difference of spatial representativeness between ASCAT and passive SM data in the
928 resampling grid cells.

929 Moreover, errors derived from SM anomalies instead of raw values were
930 considered in our analysis. The anomalies are more sensitive to capture single events
931 of drying and wetting resulting from rainfall (Dorigo et al., 2010). The coarser spatial
932 resolution of passive-based SM data can act as a lowpass filtering and may make the
933 passive SM datasets less sensitive to the drying and wetting events. By contrast, the
934 finer spatial resolution of ASCAT allows to better capture drying and wetting, and
935 this leads to a better consistency to GLDAS2. This is a plus point for the ASCAT, and
936 consequently TCA deems GLDAS2 and ASCAT SM to provide more consistent
937 observations, which may make the above phenomenon more obvious. It is worth
938 noting that ASCAT is now moving to a 6.25 km grid, and then this comparative
939 advantage of ASCAT may become even more prominent. Considering careful
940 upscaling strategy, such as studies in Albergel et al. (2010), Crow et al. (2012), and
941 Colliander et al. (2017), for ASCAT observations overlapping the GLDAS2 grids
942 may account for this problem. However, the impact of spatial resampling approaches

943 on the TCA error estimates are not reported in current literature and this point of view
944 needs further investigation in future studies. Recently, Gruber et al. (2020) also
945 proposed this concern. This result also reminds researchers to carefully process SM
946 datasets prior to TCA and prudently interpret the resulting errors in their studies.

947 It is worth mentioning that SMAPIB is a newly developed SM product and it was
948 applied in TCA only in a few studies, such as Zheng et al. (2022). Our analysis
949 suggests that TCA errors derived from SMAPIB have relatively different
950 performances compared with SMOSL3, SMOSIC, and SMAPL3 SM products, such
951 as the relatively low correlations with ground-based errors in the scatterplot of Fig. 4
952 (a), the correlation and RMSE comparisons in Fig. 5, the RD analysis in Fig. 7, and
953 the binary assessments in Fig. 9. The discrepancy between SMAPIB and other three
954 L-band SM products results from the fact that currently SMAPIB only provides SM
955 observations acquired from the descending (06:00 AM) orbits whereas other SM
956 products are reprocessed and binned into daily observations in our analysis. TCA
957 requires the inputs to describe the same physical quantity (Scipal et al. 2008a) and
958 therefore TCA deems SMAPIB SM to provide larger and distinctive error
959 characterizations when compared with other SM products.

960 The land cover analysis in Fig. 9 suggests that TCA has relatively less power to
961 efficiently characterize SM errors in croplands. The anthropogenic activities, such as
962 agricultural burning and harvesting, make it challenging to efficiently retrieve SM
963 content in croplands due to the noticeable temporal variability in canopy structure,
964 surface roughness, and diversity of crop types (Lawrence et al., 2014; Momen et al.,

965 2017; Colliander et al., 2017). This leads to uncorrected time-variant errors for SMOS
966 (Patton and Hornbuckle, 2013) and SMAP (Colliander et al., 2017) SM products, and
967 increases the difficulty of decoupling SM content from other environmental
968 information in the backscatter measurements for the ASCAT data. The imperfect
969 description of vegetation cover variability and anthropogenic activities is also a
970 concern for GLDAS2 SM data in croplands. Consequently, TCA- and ground-based
971 errors exhibit less consistency in croplands. Considering more accurate temporal
972 information of ancillary data in SM retrieval algorithms, such as irrigation and
973 harvesting activities, may reduce time-variant SM errors and improve the TCA
974 performance in the croplands.

975 Both TCA- and ground-based evaluations suggest ASCAT provides more and
976 more reliable observations as vegetation cover increases while passive-based SM has
977 more reliable observations in land cover types featured with a relatively thin
978 vegetation cover. The relatively poor performance of ASCAT data in sparsely
979 vegetated areas may be explained by the fact that ASCAT observations are sensitive
980 to subsurface scatterers during dry soil conditions and this phenomenon can extend
981 into semi-arid environments with sparse to low vegetation cover, which is recently
982 revealed by Wagner et al. (2022). The better performance of SMOS/SMAP than
983 ASCAT data regarding the time-variant errors in our analysis may be explained by the
984 fact that the L-band sensor makes SMOS/SMAP less sensitive to atmospheric effects
985 such as rainfall events (Reul et al., 2012), which is a key variable that influences time-
986 variant errors in SM products (Wu et al., 2021). However, it must be mentioned that

987 47% of the investigation pixels belong to grasslands in our study and the associated
988 ranking conclusions may not hold true when global landmass pixels are investigated.
989 Different choices in the geographical area, processing, and data screening can lead to
990 different conclusions and rankings of SM products. Besides, it is notable that SMOS
991 and SMAP SM products are up to the latest standard in this work. By contrast, this is
992 not the case for ASCAT SWI data as the Copernicus service did not invest in updating
993 the algorithm recently.

994 The robustness of applying TCA to temporally interpolated SM inputs was
995 validated and we found this interpolation has a small impact on the final TCA error
996 estimates. This is encouraging as one of the problems in the TCA time-variant scheme
997 is the limited available samples that can be used in a moving-time-window. A small
998 sample size used in TCA would underestimate the true value of random uncertainties
999 (Tsamalis, 2022) and therefore makes the resulting errors unreliable. Our results filled
1000 this gap and the temporal interpolation is recommended to be applied in the moving-
1001 window-based TCA scheme to address the limited sample issue. Moreover, Chen et al.
1002 (2018) applied the bootstrap method to guarantee reliable TCA estimates, and this
1003 may be another way to solve the sample issue in TCA.

1004 There are two limitations in this study. First, the value of $0.001 \text{ m}^3/\text{m}^3$ was
1005 selected as the threshold to distinguish the case that ASCAT and passive-based SM
1006 provide comparable errors from other possible cases. This threshold is significantly
1007 smaller than the value of $0.005 \text{ m}^3/\text{m}^3$ used in Al-Yaari et al. (2014). The evaluation
1008 results will change as the threshold value varies. Second, the depth discrepancy is

1009 considerable as satellite-based SM products generally present SM content in the
1010 topsoil layer of several centimeters whereas the GLDAS2 SM data represent SM
1011 simulations in a soil layer depth of 0-10 cm. During the ground-based evaluation,
1012 passive SM observations retrieved from L-band brightness temperature data are more
1013 consistent with the deeper sub-surface ground measurements (typically deeper than 5
1014 cm) regarding the soil layer. This may make the ground-based evaluation put more
1015 trust in the passive SM data compared with the ASCAT data.

1016 **5 Conclusions**

1017 In this study, we aimed at optimizing and validating the TCA technique with a
1018 focus on the rescaling approaches and applied it to compare time-variant errors for
1019 ASCAT and four passive-based SM products, i.e., SMOSL3, SMOSIC, SMAPL3,
1020 and SMAPIB. Based on the obtained results the following conclusions can be drawn:

1021 (1) Temporal interpolation introduces additional errors in the TCA error estimates.
1022 Nevertheless, it has a relatively small impact on the accuracy of the final TCA error
1023 estimates.

1024 (2) Rescaling technique has a great impact on the final TCA error estimates. The
1025 optimal combination strategy to implement TCA is applying TCA to SM anomalies
1026 with the TCA_Self rescaling technique. Errors derived from this optimal combination
1027 strategy achieve the highest correlation and the smallest RMSE when considering
1028 conventional ground-based errors as the benchmark.

1029 (3) Based on the optimal combination strategy, TCA-based errors are strongly
1030 correlated with ground-based errors and their Pearson's correlation coefficients are

1031 0.62, 0.72, 0.83, 0.89, and 0.93 for SMAPIB, SMAPL3, SMOSIC, SMOSL3, and
1032 ASCAT SM data, respectively. The RMSE values are typically smaller than 0.01
1033 m^3/m^3 for the above five SM products.

1034 (4) Considering and including time-variant errors in applications is necessary as time-
1035 variant errors typically deviate from time-invariant errors by a value greater than 50%.
1036 The relative magnitude between time-invariant and time-variant errors relies on the
1037 rescaling technique used in TCA. Time-invariant errors are greater than time-variant
1038 errors when SM products are rescaled against a reference dataset prior to TCA. By
1039 contrast, time-invariant errors are smaller than time-variant errors when error
1040 estimates are rescaled by the TCA_Self after the TCA implementation.

1041 (5) The evaluation performances are strongly consistent for the TCA- and ground-
1042 based methods. They provide consistent evaluation results in 74.7% (77.3%), 75.8%
1043 (79.8%), 79.6% (81.1%), and 78.6% (79.7%) of the investigation period on global
1044 average (median) for the TCA implementations with SMAPL3, SMAPIB, SMOSL3,
1045 and SMOSIC SM, respectively. However, they have two evident differences. First,
1046 TCA-based errors are mostly smaller than ground-based errors. Second, TCA
1047 generally underestimates ASCAT errors and overestimates passive-based SM errors
1048 when considering ground-based errors as the benchmark.

1049 (6) Both TCA- and ground-based methods suggest that ASCAT provides more and
1050 more reliable SM observations with smaller time-variant errors as vegetation cover
1051 increases while passive-based SM, i.e., SMAP and SMOS data here, provides more
1052 and more reliable observations with smaller time-variant errors in land cover types

1053 featured with a relatively thin vegetation cover.

1054 (7) Compared with other land cover types, TCA has relatively less power to
1055 efficiently characterize SM errors in croplands.

1056 Error characterization is crucial for correctly interpreting and efficiently using
1057 SM observations obtained from satellites, hydrological modeling, and ground
1058 measurements. Our results reveal that the selection of rescaling techniques has a great
1059 impact on the TCA error estimates, which did not receive much attention in current
1060 TCA studies. TCA can accurately evaluate SM products only when TCA is properly
1061 implemented. Our study is a strong reaffirmation of previous work in related studies
1062 that uses TCA and our study is crucial for communities who want to use TCA (and
1063 extends) for error estimation, data merging and data assimilation of satellite-based SM
1064 data.

1065

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1075

1076 **References**

- 1077 Albergel, C., Rüdiger, C., Carrer, D., Calvet, J.C., Fritz, N., Naeimi, V., Bartalis, Z.,
1078 Hasenauer, S., 2009. An evaluation of ASCAT surface soil moisture products with in-
1079 situ observations in Southwestern France. *Hydrol. Earth Syst. Sci.* 13, 115–124.
1080 <https://doi.org/10.5194/hess-13-115-2009>
- 1081 Albergel, C., Rüdiger, C., Pellarin, T., Calvet, J.C., Fritz, N., Froissard, F., Suquia, D.,
1082 Petitpa, A., Pignatelli, B., Martin, E., 2008. From near-surface to root-zone soil moisture
1083 using an exponential filter: An assessment of the method based on in-situ observations
1084 and model simulations. *Hydrol. Earth Syst. Sci.* 12, 1323–1337.
1085 <https://doi.org/10.5194/hess-12-1323-2008>
- 1086 Albergel, C., Calvet, J., Rosnay, P. De, Balsamo, G., Wagner, W., Hasenauer, S.,
1087 Naeimi, V., Martin, E., 2010. Cross-evaluation of modelled and remotely sensed
1088 surface soil moisture with in situ data in southwestern France. *Hydrol. Earth Syst. Sci.*
1089 2177–2191. <https://doi.org/10.5194/hess-14-2177-2010>
- 1090 Albergel, C., Rosnay, P. De, Gruhier, C., Muñoz-sabater, J., Hasenauer, S., Isaksen,
1091 L., Kerr, Y., Wagner, W., 2012. Evaluation of remotely sensed and modelled soil
1092 moisture products using global ground-based in situ observations. *Remote Sens.*
1093 *Environ.* 118, 215–226. <https://doi.org/10.1016/j.rse.2011.11.017>
- 1094 Al Bitar, A., Mialon, A., Kerr, Y.H., Cabot, F., Richaume, P., Jacquette, E., Quesney,
1095 A., Mahmoodi, A., Tarot, S., Parrens, M., Al-Yaari, A., Pellarin, T., Rodriguez-
1096 Fernandez, N., Wigneron, J.P., 2017. The global SMOS Level 3 daily soil moisture
1097 and brightness temperature maps. *Earth Syst. Sci. Data* 9, 293–315.
1098 <https://doi.org/10.5194/essd-9-293-2017>
- 1099 Al-Yaari, A., Wigneron, J.P., Ducharne, A., Kerr, Y.H., Wagner, W., De Lannoy, G.,
1100 Reichle, R., Al Bitar, A., Dorigo, W., Richaume, P., Mialon, A., 2014. Global-scale
1101 comparison of passive (SMOS) and active (ASCAT) satellite based microwave soil
1102 moisture retrievals with soil moisture simulations (MERRA-Land). *Remote Sens.*
1103 *Environ.* 152, 614–626. <https://doi.org/10.1016/j.rse.2014.07.013>
- 1104 An, R., Zhang, L., Wang, Z., Quayle-ballard, J.A., You, J., Shen, X., Gao, W., Huang,
1105 L., Zhao, Y., Ke, Z., 2015. Validation of the ESA CCI soil moisture product in China.
1106 *Int. J. Appl. Earth Obs. Geoinf.* 48, 28-36. <https://doi.org/10.1016/j.jag.2015.09.009>
- 1107 Bartalis, Z., Wagner, W., Naeimi, V., Hasenauer, S., Scipal, K., Bonekamp, H., Figa,
1108 J., and Anderson, C., 2007. Initial soil moisture retrievals from the METOP-A

- 1109 Advanced Scatterometer (ASCAT). *Geophys. Res. Lett.* 34, L20401,
1110 doi:10.1029/2007GL031088.
- 1111 Balsamo, G., Albergel, C., Beljaars, A., Boussetta, S., Brun, E., Cloke, H., Dee, D.,
1112 Dutra, E., Muñoz-Sabater, J., Pappenberger, F., de Rosnay, P., Stockdale, T., and
1113 Vitart, F.: ERA-Interim/Land: a global land surface reanalysis data set, *Hydrol. Earth*
1114 *Syst. Sci.*, 19, 389–407, doi:10.5194/hess-19-389-2015, 2015.
- 1115 Bell, J.E., Palecki, M.A., Baker, C.B., Collins, W.G., Lawrimore, J.H., Leeper, R.,
1116 Hall, M.E., Kochendorfer, J., Meyers, T.P., Wilson, T., et al., 2013, U.S. Climate
1117 Reference Network Soil Moisture and Temperature Observations. *J. Hydrometeorol.*
1118 14, 977–988. <https://doi.org/10.1175/JHM-D-12-0146.1>
- 1119 Biddoccu, M., Ferraris, S., Opsi, F., Cavallo, E., 2016. Long-term monitoring of soil
1120 management effects on runoff and soil erosion in sloping vineyards in Alto
1121 Monferrato (North–West Italy). *Soil Tillage Res.* 155, 176–189.
1122 <https://doi.org/10.1016/j.still.2015.07.005>
- 1123 Bircher, S., Skou, N., Jensen, K.H., Walker, J.P., Rasmussen, L., 2012. A soil
1124 moisture and temperature network for SMOS validation in Western Denmark. *Hydrol.*
1125 *Earth Syst. Sci.* 16, 1445–1463. <https://doi.org/10.5194/hess-16-1445-2012>
- 1126 Bolten, J.D., Crow, W.T., Jackson, T.J., Zhan, X., Reynolds, C.A., 2010. Evaluating
1127 the Utility of Remotely Sensed Soil Moisture Retrievals for Operational Agricultural
1128 Drought Monitoring. *IEEE J. Sel. Top. Appl. Earth Obs. Remote Sens.* 3, 57–66.
1129 <https://doi.org/10.1109/JSTARS.2009.2037163>
- 1130 Brocca, L., Melone, F., Moramarco, T., Wagner, W., Hasenauer, S., 2010. ASCAT
1131 soil wetness index validation through in situ and modeled soil moisture data in central
1132 Italy. *Remote Sens. Environ.* 114, 2745–2755.
1133 <https://doi.org/10.1016/j.rse.2010.06.009>
- 1134 Brocca, L., Hasenauer, S., Lacava, T., Melone, F., Moramarco, T., Wagner, W.,
1135 Dorigo, W., Matgen, P., Martínez-fernández, J., Llorens, P., Latron, J., Martin, C.,
1136 Bittelli, M., 2011. Soil moisture estimation through ASCAT and AMSR-E sensors:
1137 An intercomparison and validation study across Europe. *Remote Sens. Environ.* 115,
1138 3390–3408. <https://doi.org/10.1016/j.rse.2011.08.003>
- 1139 Brocca, L., Melone, F., Moramarco, T., Wagner, W., Albergel, C., 2013. Scaling and
1140 filtering approaches for the use of satellite soil moisture observations. *Remote*
1141 *Sensing of Energy Fluxes and Soil Moisture Content.* CRC Press. 411–426,
1142 <http://dx.doi.org/10.1201/b15610-21>.
- 1143 Calvet, J. -C., Fritz, N., Froissard, F., Suquia, D., Petitpa, A., & Piguet, B., 2007. In
1144 situ soil moisture observations for the CAL/VAL of SMOS: the SMOSMANIA

1145 network. IEEE Int. Geosci. Remote Sens. Symp., IGARSS, Barcelona, Spain, 1196–
1146 1199, doi:10.1109/IGARSS.2007.4423019.

1147 Cappelaere, B., Descroix, L., Lebel, T., Boulain, N., Ramier, D., Laurent, J., Favreau,
1148 G., Boubkraoui, S., Nazoumou, Y., Oi, M., Ottlé, C., Quantin, G., 2009. The AMMA-
1149 CATCH experiment in the cultivated Sahelian area of south-west Niger –
1150 Investigating water cycle response to a fluctuating climate and changing environment.
1151 J. Hydrol. 375, 34–51. <https://doi.org/10.1016/j.jhydrol.2009.06.021>

1152 Chakravorty, A., Ram, B., Prakash, O., Dhanya, C.T., 2016. A regional scale
1153 performance evaluation of SMOS and ESA-CCI soil moisture products over India
1154 with simulated soil moisture from MERRA-Land. Remote Sens. Environ. 186, 514–
1155 527. <https://doi.org/10.1016/j.rse.2016.09.011>

1156 Chen, F., Crow, W.T., Bindlish, R., Colliander, A., Burgin, M.S., Asanuma, J., Aida,
1157 K., 2018. Global-scale evaluation of SMAP, SMOS and ASCAT soil moisture
1158 products using triple collocation. Remote Sens. Environ. 214, 1–13.
1159 <https://doi.org/10.1016/j.rse.2018.05.008>

1160 Chen, F., Crow, W.T., Colliander, A., Cosh, M.H., Jackson, T., Bindlish, R., Reichle,
1161 R., Chan, S.K., Bosch, D., Starks, P., Goodrich, D., Seyfried, M., 2016. Application
1162 of triple collocation in ground-based validation of soil moisture active/passive (SMAP)
1163 level 2 data products. IEEE J. Sel. Top. Appl. Earth Obs. Remote Sens. 10 (2), 489–
1164 502. <https://doi.org/10.1109/JSTARS.2016.2569998>

1165 Chen, F., Crow, W.T., Colliander, A., Cosh, M.H., Jackson, T.J., Bindlish, R., Reichle,
1166 R.H., Chan, S.K., Bosch, D.D., Starks, P.J., Goodrich, D.C., Seyfried, M.S., 2017.
1167 Application of Triple Collocation in Ground-Based Validation of Soil Moisture
1168 Active/Passive (SMAP) Level 2 Data Products. IEEE J. Sel. Top. Appl. Earth Obs.
1169 Remote Sens. 10, 489–502. <https://doi.org/10.1109/JSTARS.2016.2569998>

1170 Colliander, A., Jackson, T.J., Bindlish, R., Chan, S., Das, N., Kim, S.B., Cosh, M.H.,
1171 Dunbar, R.S., Dang, L., Pashaian, L., Asanuma, J., Aida, K., Berg, A., Rowlandson,
1172 T., Bosch, D., Caldwell, T., Caylor, K., Goodrich, D., al Jassar, H., Lopez-Baeza, E.,
1173 Martínez-Fernández, J., González-Zamora, A., Livingston, S., McNairn, H., Pacheco,
1174 A., Moghaddam, M., Montzka, C., Notarnicola, C., Niedrist, G., Pellarin, T., Prueger,
1175 J., Pulliainen, J., Rautiainen, K., Ramos, J., Seyfried, M., Starks, P., Su, Z., Zeng, Y.,
1176 van der Velde, R., Thibeault, M., Dorigo, W., Vreugdenhil, M., Walker, J.P., Wu, X.,
1177 Monerris, A., O’Neill, P.E., Entekhabi, D., Njoku, E.G., Yueh, S., 2017. Validation of
1178 SMAP surface soil moisture products with core validation sites. Remote Sens.
1179 Environ. 191, 215–231. <https://doi.org/10.1016/j.rse.2017.01.021>

1180 Crow, W. T., A. A. Berg, M. H. Cosh, A. Loew, B. P. Mohanty, R. Panciera, P. de
1181 Rosnay, D. Ryu, and J. P. Walker, 2012. Upscaling sparse ground-based soil moisture

1182 observations for the validation of coarse-resolution satellite soil moisture products.
 1183 *Rev. Geophys.* 50, RG2002, doi:10.1029/2011RG000372.

1184 Crow, W. T., Su, C.-H., Ryu, D., and Yilmaz, M. T., 2015. Optimal averaging of soil
 1185 moisture predictions from ensemble land surface model simulations. *Water Resour.*
 1186 *Res.* 51, 9273–9289. <https://doi.org/10.1002/2015WR016944>

1187 Daly, E., Porporato, A., 2005. A review of soil moisture dynamics: From rainfall
 1188 infiltration to ecosystem response. *Environ. Eng. Sci.* 22, 9–24.
 1189 <https://doi.org/10.1089/ees.2005.22.9>

1190 Dorigo, W.A., Gruber, A., Jeu, R.A.M. De, Wagner, W., Stacke, T., Loew, A.,
 1191 Albergel, C., Brocca, L., Chung, D., Parinussa, R.M., Kidd, R., 2015. Evaluation of
 1192 the ESA CCI soil moisture product using ground-based observations. *Remote Sens.*
 1193 *Environ.* 162, 380-395. <https://doi.org/10.1016/j.rse.2014.07.023>

1194 Dorigo, W.A., Scipal, K., Parinussa, R.M., Liu, Y.Y., Wagner, W., De Jeu, R.A.M.,
 1195 Naeimi, V., 2010. Error characterisation of global active and passive microwave soil
 1196 moisture datasets. *Hydrol. Earth Syst. Sci.* 14, 2605–2616.
 1197 <https://doi.org/10.5194/hess-14-2605-2010>

1198 Dorigo, W.A., Wagner, W., Hohensinn, R., Hahn, S., Paulik, C., Xaver, A., Gruber,
 1199 A., Drusch, M., Mecklenburg, S., Van Oevelen, P., Robock, A., Jackson, T., 2011.
 1200 The International Soil Moisture Network: A data hosting facility for global in situ soil
 1201 moisture measurements. *Hydrol. Earth Syst. Sci.* 15, 1675–1698.
 1202 <https://doi.org/10.5194/hess-15-1675-2011>

1203 Dorigo, W., Xaver, A., Vreugdenhil, M., Gruber, A., A, H., Sanchis-Dufau, A.,
 1204 Zamojski, D., Cordes, C., Wagner, W., Drusch, M., 2013. Global automated quality
 1205 control of in situ soil moisture data from the international soil moisture network.
 1206 *Vadose Zone J.* 12 (3). <https://doi.org/10.2136/vzj2012.0097>

1207 Doubkova, M., Dorigo, W., Hegyiova, A., and Wagner, W., 2012. Evaluation of the
 1208 ASAR GM soil moisture product, *IEEE Int. Geosci. Remote Sens. Symp.*, 2012,
 1209 1227-1230. <https://doi.org/10.1109/IGARSS.2012.6351323>

1210 Dong, J., Wei, L., Chen, X., Duan, Z., Lu, Y., 2020. An instrument variable based
 1211 algorithm for estimating cross-correlated hydrological remote sensing errors. *J.*
 1212 *Hydrol.* 581, 124413. <https://doi.org/10.1016/j.jhydrol.2019.124413>

1213 Draper, C.S., Walker, J.P., Steinle, P.J., de Jeu, R.A.M., Holmes, T.R.H., 2009. An
 1214 evaluation of AMSR-E derived soil moisture over Australia. *Remote Sens. Environ.*
 1215 113, 703–710. <https://doi.org/10.1016/j.rse.2008.11.011>

- 1216 Draper, C., Reichle, R., Jeu, R. De, Naeimi, V., Parinussa, R., Wagner, W., 2013.
 1217 Estimating root mean square errors in remotely sensed soil moisture over continental
 1218 scale domains. *Remote Sens. Environ.* 137, 288–298.
 1219 <https://doi.org/10.1016/j.rse.2013.06.013>
- 1220 Drusch, M., Wood, E.F., Gao, H., 2005. Observation operators for the direct
 1221 assimilation of TRMM microwave imager retrieved soil moisture. *Geophys. Res. Lett.*
 1222 32, 32–35. <https://doi.org/10.1029/2005GL023623>
- 1223 Entekhabi, D., Njoku, E.G., O’Neill, P.E., Kellogg, K.H., Crow, W.T., Edelstein,
 1224 W.N., Entin, J.K., Goodman, S.D., Jackson, T.J., Johnson, J., Kimball, J., Piepmeier,
 1225 J.R., Koster, R.D., Martin, N., McDonald, K.C., Moghaddam, M., Moran, S., Reichle,
 1226 R., Shi, J.C., Spencer, M.W., Thurman, S.W., Tsang, L., Van Zyl, J., 2010. The soil
 1227 moisture active passive (SMAP) mission. *Proc. IEEE* 98, 704–716.
 1228 <https://doi.org/10.1109/JPROC.2010.2043918>
- 1229 Famiglietti, J.S., Ryu, D., Berg, A.A., Rodell, M., Jackson, T.J., 2008. Field
 1230 observations of soil moisture variability across scales. *Water Resour. Res.*, 44,
 1231 W12602, doi:10.1029/2008WR007323.
- 1232 Fascetti, F., Pierdicca, N., Pulvirenti, L., Crapolicchio, R., Muñoz-Sabater, J., 2016. A
 1233 comparison of ASCAT and SMOS soil moisture retrievals over Europe and Northern
 1234 Africa from 2010 to 2013. *Int. J. Appl. Earth Obs. Geoinf.* 45, 135–142.
 1235 <https://doi.org/10.1016/j.jag.2015.09.008>
- 1236 Friedl, M.A., McIver, D.K., Hodges, J.C.F., Zhang, X.Y., Muchoney, D., Strahler,
 1237 A.H., Woodcock, C.E., Gopal, S., Schneider, A., Cooper, A., Baccini, A., Gao, F.,
 1238 Schaaf, C., 2002. Global land cover mapping from MODIS: Algorithms and early
 1239 results. *Remote Sens. Environ.* 83, 287–302. [https://doi.org/10.1016/S0034-4257\(02\)00078-0](https://doi.org/10.1016/S0034-4257(02)00078-0)
- 1241 Gruber, A., Su, C.-H., Zwieback, S., Crow, W., Dorigo, W., Wagner, W., 2016a.
 1242 Recent advances in (soil moisture) triple collocation analysis. *Int. J. Appl. Earth Obs. Geoinf.* 45, 200–211. <https://doi.org/10.1016/j.jag.2015.09.002>
- 1244 Gruber, A., Su, C.-H., Crow, W.T., Zwieback, S., Dorigo, W. a, 2016b. Estimating
 1245 error cross-correlations in soil moisture data sets using extended collocation analysis.
 1246 *J. Geophys. Res. Atmos.* 1208–1219. <https://doi.org/10.1002/2015JD024027>
- 1247 Gruber, A., Crow, W., Dorigo, W., Wagner, W., 2015. The potential of 2D Kalman
 1248 filtering for soil moisture data assimilation. *Remote Sens. Environ.* 171, 137–148.
 1249 <https://doi.org/10.1016/j.rse.2015.10.019>

1250 Gruber, A., Dorigo, W.A., Crow, W., Wagner, W., Member, S., 2017. Triple
1251 Collocation-Based Merging of Satellite Soil Moisture Retrievals. *IEEE Trans Geosci*
1252 *Remote Sens.* 55(12), 6780–6792. <https://doi.org/10.1109/TGRS.2017.2734070>.

1253 Gruber, A., Lannoy, G. De, Crow, W., 2019. A Monte Carlo based adaptive Kalman
1254 filtering framework for soil moisture data assimilation. *Remote Sens. Environ.* 228,
1255 105–114. <https://doi.org/10.1016/j.rse.2019.04.003>

1256 Gruber, A., Paloscia, S., Santi, E., Notarnicola, C., Pasolli, L., Smolander, T.,
1257 Pulliainen, J., Mittelbach, H., Dorigo, W., Wagner, W., 2014. Performance inter-
1258 comparison of soil moisture retrieval models for the MetOp-A ASCAT instrument.
1259 *Geoscience and Remote Sensing Symposium (IGARSS)*. 2014, 2455–2458.
1260 <https://doi.org/10.1109/IGARSS.2014.6946969>

1261 Gruber, A., De Lannoy, G., Albergel, C., Al-Yaari, A., Brocca, L., Calvet, J.C.,
1262 Colliander, A., Cosh, M., Crow, W., Dorigo, W., Draper, C., Hirschi, M., Kerr, Y.,
1263 Konings, A., Lahoz, W., McColl, K., Montzka, C., Muñoz-Sabater, J., Peng, J.,
1264 Reichle, R., Richaume, P., Rüdiger, C., Scanlon, T., van der Schalie, R., Wigneron,
1265 J.P., Wagner, W., 2020. Validation practices for satellite soil moisture retrievals:
1266 What are (the) errors? *Remote Sens. Environ.* 244, 111806.
1267 <https://doi.org/10.1016/j.rse.2020.111806>

1268 Gruber, A., Dorigo, W.A., Zwieback, S., Xaver, A., Wagner, W., 2013.
1269 Characterizing Coarse-Scale Representativeness of in situ Soil Moisture
1270 Measurements from the International Soil Moisture Network. *Vadose Zo. J.* 12,
1271 vzj2012.0170. <https://doi.org/10.2136/vzj2012.0170>

1272 Houser, P.R., Shuttleworth, W.J., Famiglietti, J.S., Gupta, H. V., Syed, K.H.,
1273 Goodrich, D.C., 1998. Integration of soil moisture remote sensing and hydrologic
1274 modeling using data assimilation. *Water Resour. Res.* 34, 3405–3420.
1275 <https://doi.org/10.1029/1998WR900001>

1276 Jackson, T.J., 1993. III. Measuring surface soil moisture using passive microwave
1277 remote sensing. *Hydrol. Process.* 7, 139–152.
1278 <https://doi.org/10.1002/hyp.3360070205>

1279 Kerr, Y.H., Waldteufel, P., Richaume, P., Wigneron, J.P., Ferrazzoli, P., Mahmoodi,
1280 A., Al Bitar, A., Cabot, F., Gruhier, C., Juglea, S.E., Leroux, D., Mialon, A., Delwart,
1281 S., 2012. The SMOS soil moisture retrieval algorithm. *IEEE Trans. Geosci. Remote*
1282 *Sens.* 50, 1384–1403. <https://doi.org/10.1109/TGRS.2012.2184548>

1283 Khan, M.Z.K., Mehrotra, R., Sharma, A., Sankarasubramanian, A., 2014. Global sea
1284 surface temperature forecasts using an improved multimodel approach. *J. Clim.* 27,
1285 3505–3515. <https://doi.org/10.1175/JCLI-D-13-00486.1>

- 1286 Kim, S., Parinussa, R.M., Liu, Y.Y., Johnson, F.M., Sharma, A., 2016. Merging
1287 alternate remotely-sensed soil moisture retrievals using a non-static model
1288 combination approach. *Remote Sens.* 8, 1–16. <https://doi.org/10.3390/rs8060518>
- 1289 Kim, H., Parinussa, R., Konings, A.G., Wagner, W., Cosh, M.H., Lakshmi, V.,
1290 Zohaib, M., Choi, M., 2018. Global-scale assessment and combination of SMAP with
1291 ASCAT (active) and AMSR2 (passive) soil moisture products. *Remote Sens. Environ.*
1292 204, 260–275. <https://doi.org/10.1016/j.rse.2017.10.026>
- 1293 Kim, H., Wigneron, J., Kumar, S., Dong, J., Wagner, W., Cosh, M.H., Bosch, D.D.,
1294 Holifield, C., Starks, P.J., Seyfried, M., Sol, I., Atmosphère, P., Recherche, U.M. De,
1295 National, I., Recherche, D., Inrae, A., 2020. Global scale error assessments of soil
1296 moisture estimates from microwave-based active and passive satellites and land
1297 surface models over forest and mixed irrigated/dryland agriculture regions. *Remote*
1298 *Sens. Environ.* 251, 112052. <https://doi.org/10.1016/j.rse.2020.112052>
- 1299 Kim, S., Sharma, A., Liu, Y.Y., Young, S.I., 2022. Rethinking Satellite Data Merging:
1300 From Averaging to SNR Optimization. *IEEE Trans. Geosci. Remote Sens.* 60.
1301 <https://doi.org/10.1109/TGRS.2021.3107028>
- 1302 Kim, S., Dong, J., Sharma, A., 2021. A Triple Collocation-Based Comparison of
1303 Three L-Band Soil Moisture Datasets, SMAP, SMOS-IC, and SMOS, Over Varied
1304 Climates and Land Covers. *Front. Water* 3, 1–12.
1305 <https://doi.org/10.3389/frwa.2021.693172>
- 1306 Larson, K. M., Small, E. E., Gutmann, E. D., Bilich, A. L., Braun, J. J., & Zavorotny,
1307 V. U., 2008. Use of GPS receivers as a soil moisture network for water cycle studies.
1308 *Geophysical Research Letters*, 35, 24. <https://doi.org/10.1029/2008GL036013>
- 1309 Lawrence, H., Wigneron, J.P., Richaume, P., Novello, N., Grant, J., Mialon, A., Al
1310 Bitar, A., Merlin, O., Guyon, D., Leroux, D., Bircher, S., Kerr, Y., 2014. Comparison
1311 between SMOS Vegetation Optical Depth products and MODIS vegetation indices
1312 over crop zones of the USA. *Remote Sens. Environ.* 140, 396–406.
1313 <https://doi.org/10.1016/j.rse.2013.07.021>
- 1314 Leroux, D.J., Kerr, Y.H., Richaume, P., Fieuzal, R., 2013. Spatial distribution and
1315 possible sources of SMOS errors at the global scale. *Remote Sens. Environ.* 133, 240–
1316 250. <https://doi.org/10.1016/j.rse.2013.02.017>
- 1317 Li, X., Wigneron, J.P., Fan, L., Frappart, F., Yueh, S.H., Colliander, A., Ebtehaj, A.,
1318 Gao, L., Fernandez-Moran, R., Liu, X., Wang, M., Ma, H., Moisy, C., Ciais, P., 2022.
1319 A new SMAP soil moisture and vegetation optical depth product (SMAP-IB):
1320 Algorithm, assessment and inter-comparison. *Remote Sens. Environ.* 271, 112921.
1321 <https://doi.org/10.1016/j.rse.2022.112921>

- 1322 Li, H., Chai, L., Crow, W., Dong, J., Liu, S., Zhao, S., 2022. The reliability of
 1323 categorical triple collocation for evaluating soil freeze/thaw datasets. *Remote Sens.*
 1324 *Environ.* 281, 113240. <https://doi.org/10.1016/j.rse.2022.113240>
- 1325 Liu, J., Chai, L., Dong, J., Zheng, D., Wigneron, J.P., Liu, S., Zhou, J., Xu, T., Yang,
 1326 S., Song, Y., Qu, Y., Lu, Z., 2021. Uncertainty analysis of eleven multisource soil
 1327 moisture products in the third pole environment based on the three-corned hat method.
 1328 *Remote Sens. Environ.* 255, 112225. <https://doi.org/10.1016/j.rse.2020.112225>
- 1329 Loew, A., Schlenz, F., 2011. A dynamic approach for evaluating coarse scale satellite
 1330 soil moisture products. *Hydrol. Earth Syst. Sci.* 15, 75–90.
 1331 <https://doi.org/10.5194/hess-15-75-2011>
- 1332 Mattar, C., Santamaría-Artigues, A., Durán-Alarcón, C., Olivera-Guerra, L., Fuster,
 1333 R., Borvarán, D., 2016. The LAB-net soil moisture network: application to thermal
 1334 remote sensing and surface energy balance. *Data* 1, 6.
 1335 <https://doi.org/10.3390/data1010006>
- 1336 Miralles, D.G., Crow, W.T., Cosh, M.H., 2010. Estimating spatial sampling errors in
 1337 coarse-scale soil moisture estimates derived from point-scale observations. *J.*
 1338 *Hydrometeorol.* 11, 1423–1429. <https://doi.org/10.1175/2010JHM1285.1>
- 1339 Miyaoka, K., Gruber, A., Ticconi, F., Hahn, S., Wagner, W., Figa-Saldana, J.,
 1340 Anderson, C., 2017. Triple Collocation Analysis of Soil Moisture from Metop-A
 1341 ASCAT and SMOS Against JRA-55 and ERA-Interim. *IEEE J. Sel. Top. Appl. Earth*
 1342 *Obs. Remote Sens.* 10, 2274–2284. <https://doi.org/10.1109/JSTARS.2016.2632306>
- 1343 Moghaddam, M., Entekhabi, D., Member, S., Goykhman, Y., Li, K., Liu, M., Member,
 1344 S., Mahajan, A., Nayyar, A., Shuman, D., 2010. A Wireless Soil Moisture Smart
 1345 Sensor Web Using Physics-Based Optimal Control: Concept and Initial
 1346 Demonstrations. *IEEE J. Sel. Top. Appl. Earth Obs. Remote Sens.* 3, 522–535.
 1347 <https://doi.org/10.1109/JSTARS.2010.2052918>
- 1348 Mougouin, E., Hiernaux, P., Kergoat, L., Grippa, M., Rosnay, P. De, Timouk, F., Dantec,
 1349 V. Le, Demarez, V., Lavenu, F., Arjounin, M., Lebel, T., Soumaguel, N., Ceschia, E.,
 1350 Mougouin, B., Baup, F., Frappart, F., Frison, P.L., Gardelle, J., Gruhier, C., Jarlan, L.,
 1351 Mangiarotti, S., Sanou, B., Tracol, Y., Guichard, F., Serça, D., Galy-lacaux, C.,
 1352 Seghieri, J., Becerra, S., Dia, H., Gangneron, F., Mazzega, P., 2009. The AMMA-
 1353 CATCH Gourma observatory site in Mali: Relating climatic variations to changes in
 1354 vegetation, surface hydrology, fluxes and natural resources. *J. Hydrol.* 375, 14–33.
 1355 <https://doi.org/10.1016/j.jhydrol.2009.06.045>
- 1356 Momen, M., Wood, J.D., Novick, K.A., Pangle, R., Pockman, W.T., McDowell, N.G.,
 1357 Konings, A.G., 2017. Interacting Effects of Leaf Water Potential and Biomass on

- 1358 Vegetation Optical Depth. *J. Geophys. Res. Biogeosciences* 122, 3031–3046.
1359 <https://doi.org/10.1002/2017JG004145>
- 1360 Molero, B., Leroux, D.J., Richaume, P., Kerr, Y.H., Merlin, O., Cosh, M.H., Bindlish,
1361 R., 2018. Multi-Timescale Analysis of the Spatial Representativeness of In Situ Soil
1362 Moisture Data within Satellite Footprints. *J. Geophys. Res. Atmos.* 123, 3–21.
1363 <https://doi.org/10.1002/2017JD027478>
- 1364 Narasimhan, B., Srinivasan, R., 2005. Development and evaluation of Soil Moisture
1365 Deficit Index (SMDI) and Evapotranspiration Deficit Index (ETDI) for agricultural
1366 drought monitoring. *Agric. For. Meteorol.* 133, 69–88.
1367 <https://doi.org/10.1016/j.agrformet.2005.07.012>
- 1368 Osenga, E.C., Arnott, J.C., Endsley, K.A., Katzenberger, J.W., 2019. Bioclimatic and
1369 soil moisture monitoring across elevation in a mountain watershed: opportunities for
1370 research and resource management. *Water Resour. Res.* 55, 2493– 2503.
1371 <https://doi.org/10.1029/2018WR023653>
- 1372 O'Neill, P. E., S. Chan, E. G. Njoku, T. Jackson, R. Bindlish, and J. Chaubell. (2021).
1373 SMAP L3 Radiometer Global Daily 36 km EASE-Grid Soil Moisture, Version 8
1374 [Data Set]. Boulder, Colorado USA. NASA National Snow and Ice Data Center
1375 Distributed Active Archive Center. <https://doi.org/10.5067/OMHVSARGFX380>. Date
1376 Accessed 07-20-2022.
- 1377 Paulik, C., Dorigo, W., Wagner, W., Kidd, R., 2014. Validation of the ASCAT soil
1378 water index using in situ data from the International Soil moisture network. *Int. J.*
1379 *Appl. Earth Obs. Geoinf.* 30, 1–8. <https://doi.org/10.1016/j.jag.2014.01.007>
- 1380 Patton, J., Hornbuckle, B., 2013. Initial validation of SMOS vegetation optical
1381 thickness in Iowa. *IEEE Geosci. Remote Sens. Lett.* 10, 3791–3794.
1382 <https://doi.org/10.1109/IGARSS.2012.6350491>
- 1383 Pellarin, T., Laurent, J.P., Cappelaere, B., Decharme, B., Descroix, L., Ramier, D.,
1384 2009. Hydrological modelling and associated microwave emission of a semi-arid
1385 region in South-western Niger. *J. Hydrol.* 375, 262–272.
1386 <https://doi.org/10.1016/j.jhydrol.2008.12.003>
- 1387 Peng, J., Tanguy, M., Robinson, E.L., Pinnington, E., Evans, J., Ellis, R., Cooper, E.,
1388 Hannaford, J., Blyth, E., Dadson, S., 2021. Estimation and evaluation of high-
1389 resolution soil moisture from merged model and Earth observation data in the Great
1390 Britain. *Remote Sens. Environ.* 264, 112610.
1391 <https://doi.org/10.1016/j.rse.2021.112610>
- 1392 Pierdicca, N., Fascetti, F., Pulvirenti, L., Crapolicchio, R., Muñoz-sabater, J., 2015.
1393 Analysis of ASCAT, SMOS, in-situ and land model soil moisture as a regionalized

- 1394 variable over Europe and North Africa. *Remote Sens. Environ.* 170, 280–289.
1395 <https://doi.org/10.1016/j.rse.2015.09.005>
- 1396 Pierdicca, N., Fascetti, F., Pulvirenti, L., Crapolicchio, R., 2017. Error
1397 Characterization of Soil Moisture Satellite Products: Retrieving Error Cross-
1398 Correlation Through Extended Quadruple Collocation. *IEEE J. Sel. Top. Appl. Earth
1399 Obs. Remote Sens.* 10, 4522–4530. <https://doi.org/10.1109/JSTARS.2017.2714025>
- 1400 Reichle, R.H., Koster, R.D., 2004. Bias reduction in short records of satellite soil
1401 moisture. *Geophys. Res. Lett.* 31, L19501. <https://doi.org/10.1029/2004GL020938>
- 1402 Reichle, R.H., 2008. Data assimilation methods in the Earth sciences. *Adv. Water
1403 Resour.* 31, 1411–1418. <https://doi.org/10.1016/j.advwatres.2008.01.001>
- 1404 Rodell, M., Houser, P.R., Jambor, U., Gottschalck, J., Mitchell, K., Meng, C.J.,
1405 Arsenault, K., Cosgrove, B., Radakovich, J., Bosilovich, M., Entin, J.K., Walker, J.P.,
1406 Lohmann, D., Toll, D., 2004. The Global Land Data Assimilation System. *Bull. Am.
1407 Meteorol. Soc.* 85, 381–394. <https://doi.org/10.1175/BAMS-85-3-381>
- 1408 Rosnay, P. De, Gruhier, C., Timouk, F., Baup, F., Mougin, E., Hiernaux, P., Kergoat,
1409 L., Ledantec, V., 2009. Multi-scale soil moisture measurements at the Gourma meso-
1410 scale site in Mali. *J. Hydrol.* 375, 241–252.
1411 <https://doi.org/10.1016/j.jhydrol.2009.01.015>
- 1412 Rüdiger, C., Calvet, J., Gruhier, C., Holmes, T. R. H., de Jeu, R. A. M., & Wagner,
1413 W., 2009. An Intercomparison of ERS-Scat and AMSR-E Soil Moisture Observations
1414 with Model Simulations over France. *J. Hydrometeorol.* 10(2), 431-447.
- 1415 Reul, N., Tenerelli, J., Chapron, B., Vandemark, D., Quilfen, Y., and Kerr, Y., 2012.
1416 SMOS satellite L-band radiometer: A new capability for ocean surface remote sensing
1417 in hurricanes. *J. Geophys. Res.* 117, C02006, doi:10.1029/2011JC007474.
- 1418 Scipal, K., Holmes, T., De Jeu, R., Naeimi, V., Wagner, W., 2008a. A possible
1419 solution for the problem of estimating the error structure of global soil moisture data
1420 sets. *Geophys. Res. Lett.* 35, 2–5. <https://doi.org/10.1029/2008GL035599>
- 1421 Scipal, K., Drusch, M., Wagner, W., 2008b. Assimilation of a ERS scatterometer
1422 derived soil moisture index in the ECMWF numerical weather prediction system. *Adv.
1423 Water Resour.* 31, 1101–1112. <https://doi.org/10.1016/j.advwatres.2008.04.013>.
- 1424 Scipal, K., Scheffler, C., Wagner, W., 2005. Soil moisture-runoff relation at the
1425 catchment scale as observed with coarse resolution microwave remote sensing.
1426 *Hydrol. Earth Syst. Sci.* 9, 173–183.

- 1427 Scipal, K., Dorigo, W., De Jeu, R., 2010. Triple collocation - A new tool to determine
1428 the error structure of global soil moisture products. *Int. Geosci. Remote Sens. Symp.* 3,
1429 4426–4429. <https://doi.org/10.1109/IGARSS.2010.5652128>.
- 1430 Sims, A.P., Niyogi, D.D.S., Raman, S., 2002. Adopting drought indices for estimating
1431 soil moisture: A North Carolina case study. *Geophys. Res. Lett.* 29, 24-1-24-4.
1432 <https://doi.org/10.1029/2001GL013343>
- 1433 Smith, A.B., Walker, J.P., Western, A.W., Young, R.I., Ellett, K.M., Pipunic, R.C.,
1434 Grayson, R.B., Siriwardena, L., Chiew, F.H.S., Richter, H., 2012. The Murrumbidgee
1435 Soil Moisture Monitoring Network data set. *Water Resour. Res.* 48, 1–6.
1436 <https://doi.org/10.1029/2012WR011976>.
- 1437 Stoffelen, A., 1998. Toward the true near-surface wind speed: Error modeling and
1438 calibration using triple collocation. *J. Geophys. Res. C Ocean.* 103, 7755–7766.
1439 <https://doi.org/10.1029/97jc03180>
- 1440 Su, C., Ryu, D., Crow, W.T., Western, A.W., 2014a. Stand-alone error
1441 characterisation of microwave satellite soil moisture using a Fourier method. *Remote*
1442 *Sens. Environ.* 154, 115–126. <https://doi.org/10.1016/j.rse.2014.08.014>
- 1443 Su, C.H., Ryu, D., Crow, W.T., Western, A.W., 2014b. Beyond triple collocation:
1444 Applications to soil moisture monitoring. *J. Geophys. Res.* 119, 6419–6439.
1445 <https://doi.org/10.1002/2013JD021043>
- 1446 Su, C., Ryu, D., Young, R.I., Western, A.W., Wagner, W., 2013. Inter-comparison of
1447 microwave satellite soil moisture retrievals over the Murrumbidgee Basin, southeast
1448 Australia. *Remote Sens. Environ.* 134, 1–11. <https://doi.org/10.1016/j.rse.2013.02.016>
- 1449 Tagesson, T., Fensholt, R., Guiro, I., Rasmussen, M.O., Huber, S., Mbow, C., Garcia,
1450 M., Horion, S., Sandholt, I., Holm-Rasmussen, B., et al., 2015. Ecosystem properties
1451 of semiarid savanna grassland in West Africa and its relationship with environmental
1452 variability. *Glob. Chang. Biol.* 21, 250–264. <https://doi.org/10.1111/gcb.12734>
- 1453 Tsamalis, C., 2022. Clarifications on the equations and the sample number in triple
1454 collocation analysis using SST observations. *Remote Sens. Environ.* 272, 112936.
1455 <https://doi.org/10.1016/j.rse.2022.112936>
- 1456 Tugrul Yilmaz, M., Crow, W.T., 2013. The optimality of potential rescaling
1457 approaches in land data assimilation. *J. Hydrometeorol.* 14, 650–660.
1458 <https://doi.org/10.1175/JHM-D-12-052.1>
- 1459 Tugrul Yilmaz, M., Crow, W.T., 2014. Evaluation of assumptions in soil moisture
1460 triple collocation analysis. *J. Hydrometeorol.* 15, 1293–1302.
1461 <https://doi.org/10.1175/JHM-D-13-0158.1>

- 1462 Ulaby, F.T., Dobson, M.C., Brunfeldt, D.R., 1983. Improvement of Moisture
1463 Estimation Accuracy of Vegetation-Covered Soil by Combined Active/Passive
1464 Microwave Remote Sensing. *IEEE Trans. Geosci. Remote Sens.* GE-21, 300–307.
1465 <https://doi.org/10.1109/TGRS.1983.350557>
- 1466 Wagner, W., Lemoine, G., Rott, H., 1999. A method for estimating soil moisture from
1467 ERS Scatterometer and soil data. *Remote Sens. Environ.* 70, 191–207.
1468 [https://doi.org/10.1016/S0034-4257\(99\)00036-X](https://doi.org/10.1016/S0034-4257(99)00036-X)
- 1469 Wagner, W., Lindorfer, R., Melzer, T., Hahn, S., Bauer-marschallinger, B., Morrison,
1470 K., Calvet, J., Hobbs, S., Quast, R., Greimeister-pfeil, I., Vreugdenhil, M., 2022.
1471 Widespread occurrence of anomalous C-band backscatter signals in arid environments
1472 caused by subsurface scattering. *Remote Sens. Environ.* 276, 113025.
1473 <https://doi.org/10.1016/j.rse.2022.113025>
- 1474 Wang, X., Lü, H., Crow, W.T., Zhu, Y., Wang, Q., Su, J., Zheng, J., Gou, Q., 2021.
1475 Assessment of SMOS and SMAP soil moisture products against new estimates
1476 combining physical model , a statistical model , and in-situ observations : A case
1477 study over the Huai River Basin , China. *J. Hydrol.* 598, 126468.
1478 <https://doi.org/10.1016/j.jhydrol.2021.126468>
- 1479 Wanders, N., Karssenberg, D., De Roo, A., De Jong, S.M., Bierkens, M.F.P., 2014.
1480 The suitability of remotely sensed soil moisture for improving operational flood
1481 forecasting. *Hydrol. Earth Syst. Sci.* 18, 2343–2357. [https://doi.org/10.5194/hess-18-](https://doi.org/10.5194/hess-18-2343-2014)
1482 [2343-2014](https://doi.org/10.5194/hess-18-2343-2014)
- 1483 Western, A.W., Grayson, R.B., Blöschl, G., 2002. Scaling of Soil Moisture: A
1484 Hydrologic Perspective. *Annu. Rev. Earth Planet. Sci.* 30, 149–180.
1485 <https://doi.org/10.1146/annurev.earth.30.091201.140434>
- 1486 Wigneron, J.P., Li, X., Frappart, F., Fan, L., Al-Yaari, A., De Lannoy, G., Liu, X.,
1487 Wang, M., Le Masson, E., Moisy, C., 2021. SMOS-IC data record of soil moisture
1488 and L-VOD: Historical development, applications and perspectives. *Remote Sens.*
1489 *Environ.* 254. <https://doi.org/10.1016/j.rse.2020.112238>
- 1490 Wu, K., Ryu, D., Nie, L., Shu, H., 2021. Time-variant error characterization of SMAP
1491 and ASCAT soil moisture using Triple Collocation Analysis. *Remote Sens. Environ.*
1492 256, 112324. <https://doi.org/10.1016/j.rse.2021.112324>
- 1493 Wu, K., Shu, H., Nie, L., Jiao, Z., 2018. Triple collocation-based estimation of
1494 spatially correlated observation error covariance in remote sensing soil moisture data
1495 assimilation. *J. Appl. Remote Sens.* 12, 1–19. <https://doi.org/10.1117/1.JRS.12>

- 1496 Wu, K., Nie, L., Shu, H., 2020. A comparison of SMAP and SMOS L-band
1497 brightness temperature observations over the global landmass. *Int. J. Remote Sens.* 41,
1498 399–419. <https://doi.org/10.1080/01431161.2019.1641759>
- 1499 Wu, X., Xiao, Q., Wen, J., You, D., 2019. Direct Comparison and Triple Collocation:
1500 Which Is More Reliable in the Validation of Coarse-Scale Satellite Surface Albedo
1501 Products. *J. Geophys. Res. Atmos.* 124, 5198–5213.
1502 <https://doi.org/10.1029/2018JD029937>
- 1503 Xu, L., Chen, N., Zhang, X., Moradkhani, H., Zhang, C., Hu, C., 2021. In-situ and
1504 triple-collocation based evaluations of eight global root zone soil moisture products.
1505 *Remote Sens. Environ.* 254, 112248. <https://doi.org/10.1016/j.rse.2020.112248>
- 1506 Yang, K., Qin, J., Zhao, L., Chen, Y., Tang, W., Han, M., Chen, Z., Lv, N., Ding, B.,
1507 Wu, H., & Lin, C., 2013. A Multiscale Soil Moisture and Freeze–Thaw Monitoring
1508 Network on the Third Pole. *Bull. Amer. Meteor. Soc.* 94(12), 1907–1916.
1509 <https://doi.org/10.1175/BAMS-D-12-00203.1>
- 1510 Zacharias, S., Bogen, H., Samaniego, L., Mauder, M., Fuß, R., Pütz, T., Frenzel, M.,
1511 Schwank, M., Baessler, C., Butterbach-Bahl, K., Bens, O., Borg, E., Brauer, A.,
1512 Dietrich, P., Hajsek, I., Helle, G., Kiese, R., Kunstmann, H., Klotz, S., Munch, J.C.,
1513 Papen, H., Priesack, E., Schmid, H.P., Steinbrecher, R., Rosenbaum, U., Teutsch, G.,
1514 Vereecken, H., 2011. A Network of Terrestrial Environmental Observatories in
1515 Germany. *Vadose Zo. J.* 10, 955–973. <https://doi.org/10.2136/vzj2010.0139>.
- 1516 Zhang, R., Kim, S., Sharma, A., Lakshmi, V., 2021. Identifying relative strengths of
1517 SMAP, SMOS-IC, and ASCAT to capture temporal variability. *Remote Sens. Environ.*
1518 252, 112126. <https://doi.org/10.1016/j.rse.2020.112126>
- 1519 Zheng, J., Zhao, T., Lü, H., Shi, J., Cosh, M.H., Ji, D., Jiang, L., Cui, Q., Lu, H.,
1520 Yang, K., Wigneron, J.P., Li, X., Zhu, Y., Hu, L., Peng, Z., Zeng, Y., Wang, X., Kang,
1521 C.S., 2022. Assessment of 24 soil moisture datasets using a new in situ network in the
1522 Shandian River Basin of China. *Remote Sens. Environ.* 271.
1523 <https://doi.org/10.1016/j.rse.2022.112891>
- 1524 Zhuang, R., Zeng, Y., Manfreda, S., Su, Z., 2020. Quantifying long-term land surface
1525 and root zone soil moisture over Tibetan plateau. *Remote Sens.* 12, 509.
1526 <https://doi.org/10.3390/rs12030509>
- 1527 Zreda, M., Shuttleworth, W. J., Zeng, X., Zweck, C., Desilets, D., Franz, T., and
1528 Rosolem, R., COSMOS: the COsmic-ray Soil Moisture Observing System. *Hydrol.*
1529 *Earth Syst. Sci.*, 16, 4079–4099, <https://doi.org/10.5194/hess-16-4079-2012>, 2012.
- 1530 Zwieback, S., Colliander, A., Cosh, M.H., Martínez-Fernández, J., McNairn, H.,
1531 Starks, P.J., Thibeault, M., Berg, A., 2018. Estimating time-dependent vegetation

1532 biases in the SMAP soil moisture product. *Hydrol. Earth Syst. Sci.* 22, 4473–4489.
1533 <https://doi.org/10.5194/hess-22-4473-2018>

1534 Zwieback, S., Dorigo, W., Wagner, W., 2012. Temporal error variability of coarse
1535 scale soil moisture products - Case study in central Spain. *Int. Geosci. Remote Sens.*
1536 *Symp.* 722–725. <https://doi.org/10.1109/IGARSS.2012.6351463>

1537

1538 **Supplementary**

1539 Matlab codes with an example to estimate time-variant errors based on the TCA_Self
1540 rescaling technique is provided along with this manuscript. Please cite this manuscript
1541 if you intend to use the provided Matlab codes.

1542

1543

List of Figure Captions

1544 **Fig. 1** Spatial distribution of 759 grid cells that include selected ISMN stations.

1545 **Fig. 2** Boxplots of correlation coefficients (red) and RMSD values (blue) between time-
1546 variant errors derived from the original SM data without interpolation and the SM time series
1547 resampled by different percentage values of the original data. (a-d) exhibits the results of
1548 applying interpolation to SMOSL3, SMAPL3, SMOSIC, and ASCAT SM data, respectively.
1549 The x-axis denotes the percentage values of the original data considered in the resampling.
1550 The y-axes on the left (red color) and right (blue color) describe correlation coefficients and
1551 RMSD, respectively. The ‘N’ above each subfigure denotes the number of available samples
1552 included in the corresponding boxplot.

1553 **Fig. 3** Boxplots of the Pearson’s correlation coefficients (red) and RMSE values (blue)
1554 between ground- and TCA-based time-variant errors for the TCA implementations that
1555 consider (a) SMOSL3, (b) SMOSIC, (c) SMAPL3, (d) SMAPIB, and (e) ASCAT as the triplet
1556 inputs. In each subfigure, results derived from absolutes and anomalies are shown in the white
1557 and grey areas, respectively. The x-axis is the multiple rescaling techniques considered in the
1558 TCA. The y-axes on the left (red color) and right (blue color) describe correlation coefficients
1559 and RMSE, respectively.

1560 **Fig. 4** Scatterplots of ground- and TCA-based time-variant errors for (a) SMAPIB, (b)
1561 SMAPL3, (c) SMOSIC, (d) SMOSL3, and (e) ASCAT SM data. The original scatter points
1562 are binned in the increase of x-axis for ground-based error and y-axis for TCA-based error,
1563 respectively, and colored based on the number of points included in the bin. All the p values
1564 passed a t-test ($\alpha < 0.05$). N describes the number of experimental grid cells considered in
1565 each scatterplot.

1566 **Fig. 5** Boxplots of (a) correlation coefficients and (b) RMSE between TCA- and ground-
1567 based errors classified by six land cover types for ASCAT and multiple passive SM products.
1568 Boxplots with limited available samples (< 30) are not considered in such comparisons.

1569 **Fig. 6** Mean and median of the ORD values derived from the NORM, VAR, CDF, TCA_Self,
1570 and QCA_Self rescaling techniques for SMAPL3, SMAPIB, SMOSL3, SMOSIC, and
1571 ASCAT SM datasets. (a) and (b) represent ORD values obtained from the TCA- and ground-

1572 based evaluations, respectively.

1573 **Fig. 7** Fractions of the RD values that are greater (red) or smaller (blue) than 0 for the
1574 combination of five SM datasets, i.e., ASCAT, SMOSIC, SMOSL3, SMAPIB, and SMAPL3,
1575 and five rescaling techniques, i.e., QCA_Self, TCA_Self, CDF, VAR, and NORM. (a) and (b)
1576 present associated results based on the RD values derived from the TCA- and ground-based
1577 evaluations, respectively.

1578 **Fig. 8** Boxplots of a binary assessment of the evaluation consistency between TCA- and
1579 ground-based methods for ASCAT and four passive-based SM time-variant errors. The four
1580 columns illustrated in this figure denote the four TCA implementations that consider
1581 [GLDAS2, ASCAT, SMAPL3], [GLDAS2, ASCAT, SMAPIB], [GLDAS2, ASCAT,
1582 SMOSL3], and [GLDAS2, ASCAT, SMOSIC] as the triplet.

1583 **Fig. 9** Boxplots of the percentage values classified by six land cover types for the three cases:
1584 ASCAT (green boxes) or passive SM (blue boxes) has the smallest time-variant error and the
1585 difference between their errors is smaller than $0.001 \text{ m}^3/\text{m}^3$ (yellow boxes). (a-d) represent the
1586 TCA- (white areas) and ground-based (grey areas) assessments that consider SMOSL3,
1587 SMOSIC, SMAPL3, and SMAPIB SM data, respectively. To guarantee reliable results,
1588 boxplots associated with limited available samples (< 30) are excluded. The two pies below
1589 each subfigure exhibit the overall performance of the three cases that include all available
1590 samples for the TCA- (on the left) and ground-based (on the right) assessments.

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List of Table Captions

1593 **Table 1** A summary of the SM products used in this work

1594 **Table 2** ISMN ground stations from sparse networks

1595 **Table 3** Pearson's correlation coefficients between ground- and TCA-based time-invariant errors

1596 for SMOSL3, SMOSIC, SMAPL3, SMAPIB, and ASCAT SM products. The time-invariant errors

1597 are derived from multiple rescaling techniques. The first and second recommendation strategies

1598 (excluding QCA_Self) to implement TCA are highlighted with green and yellow colors,

1599 respectively. All the ρ values passed a t-test ($\alpha < 0.05$)

1600 **Table 4** The same as Table 3 but for the RMSE values